

Three-dimensional imaging of the Martian polar ice caps from orbit with the MRO Shallow Radar sounder

The Shallow Radar (SHARAD) on
the Mars Reconnaissance Orbiter



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with

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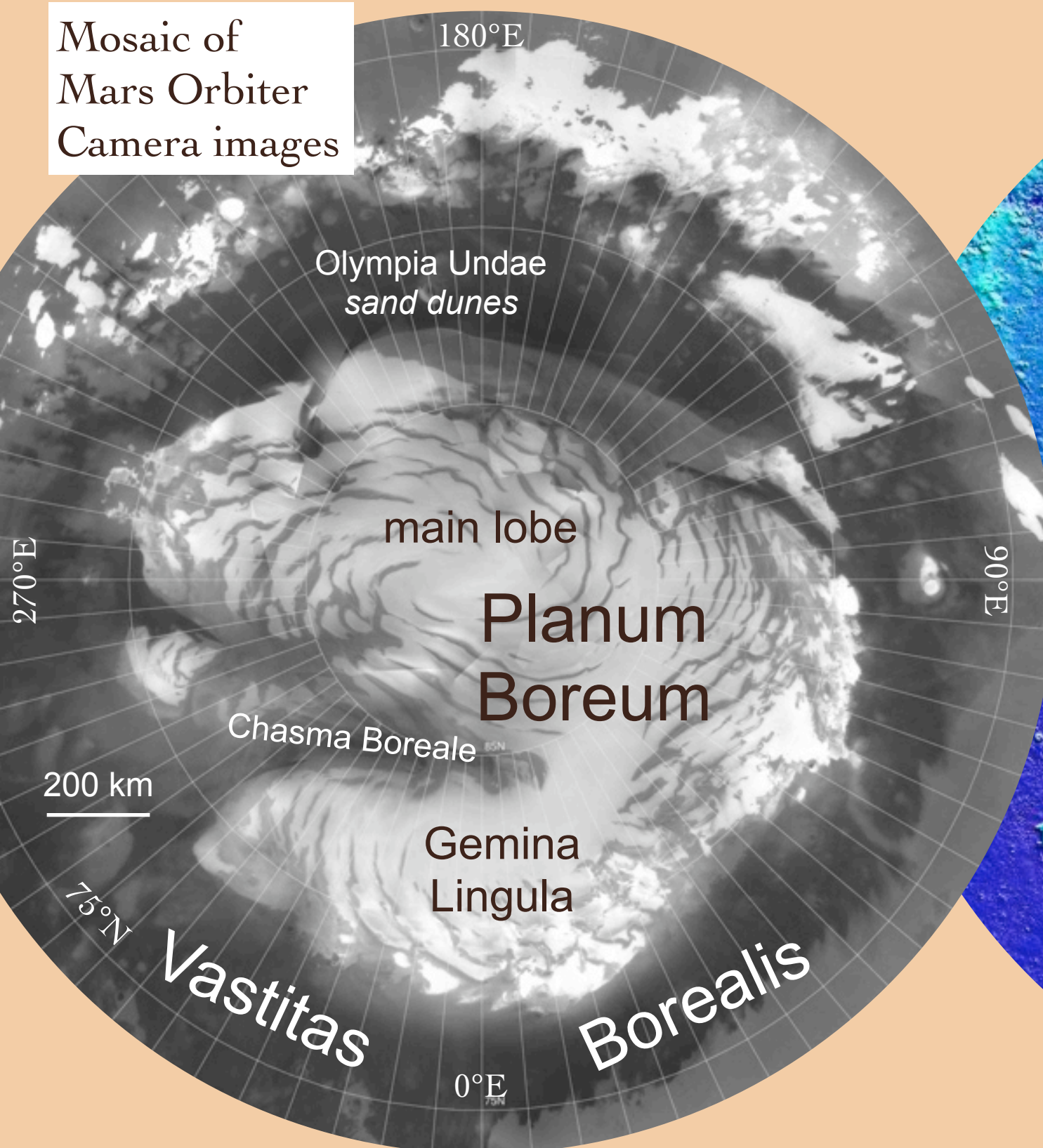
Outline

- Overview of Martian north polar region and the SHARAD instrument.
- Summary of polar science results from the collection of 2-D radargrams.
- First results for 3-D processing of the north polar data volume.
- Future plans and Conclusions.

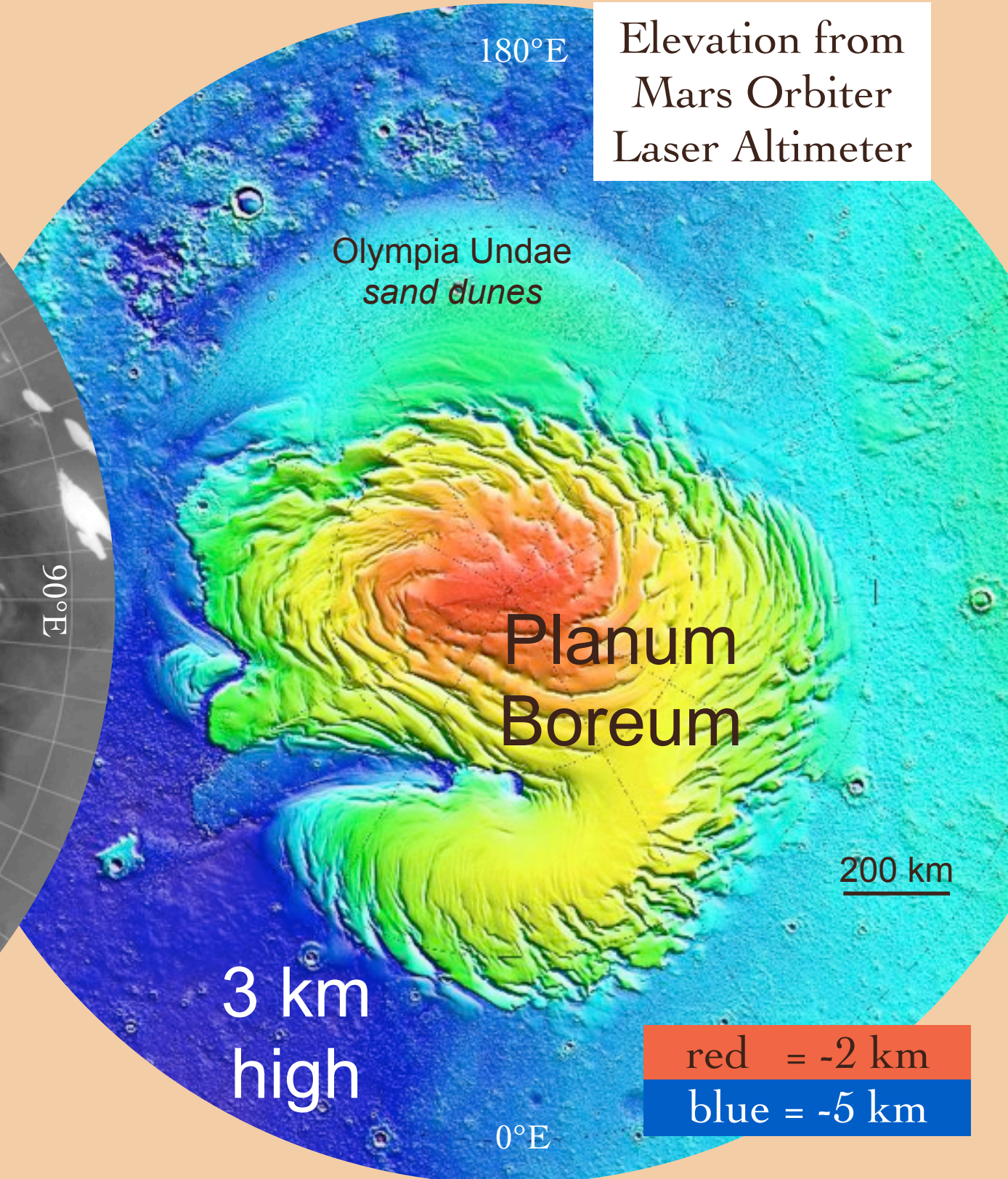
The North Polar Region of Mars

Bright, high-standing ice-rich deposits cut by layer-exposing troughs and surrounded by basaltic dunes.

Mosaic of
Mars Orbiter
Camera images



Elevation from
Mars Orbiter
Laser Altimeter

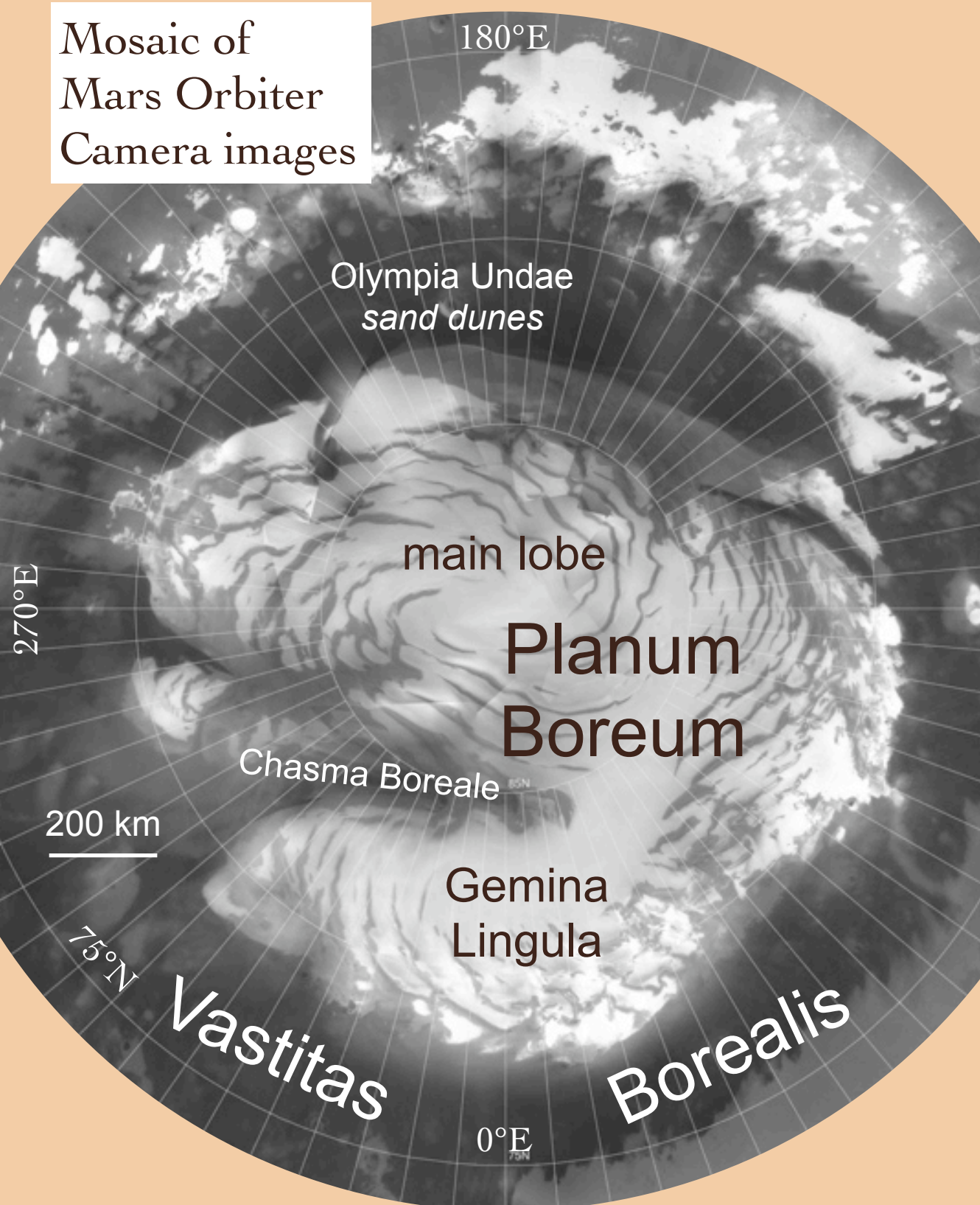


red = -2 km
blue = -5 km

The North Polar Region of Mars

Bright, high-standing ice-rich deposits cut by layer-exposing troughs and surrounded by basaltic dunes.

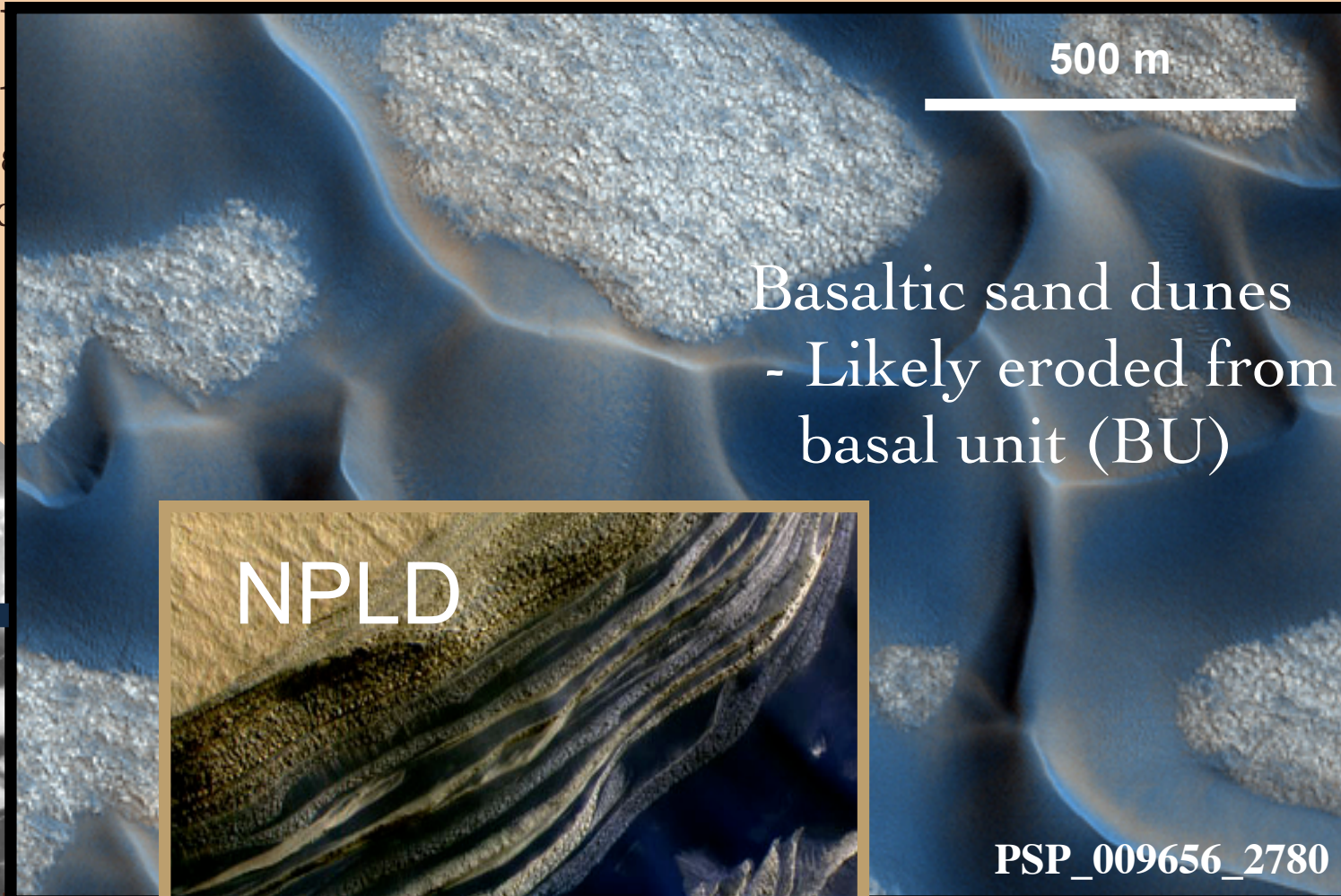
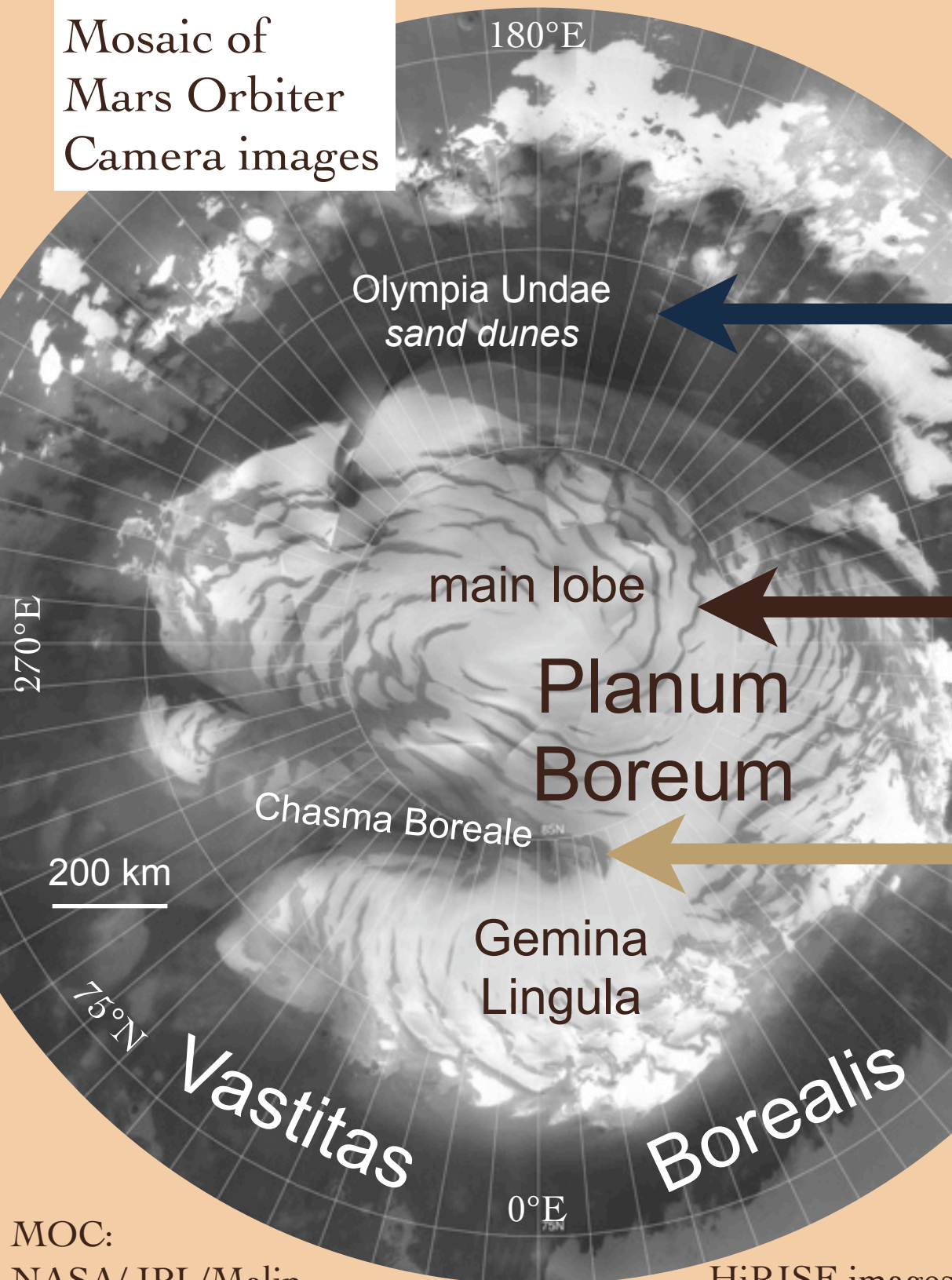
Mosaic of
Mars Orbiter
Camera images



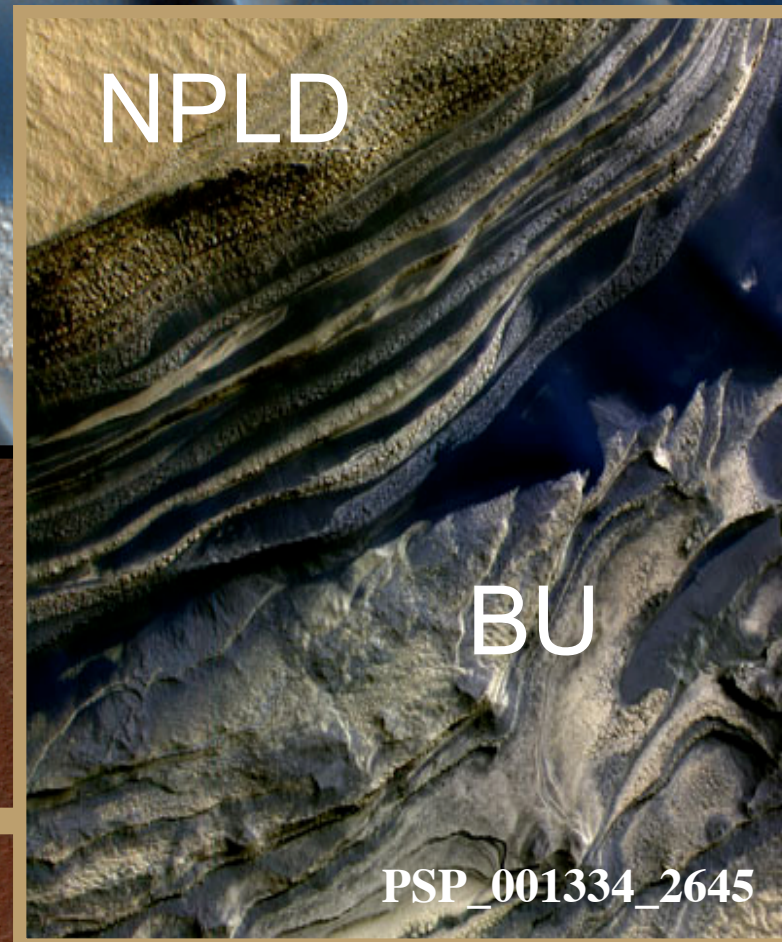
The North Pole

Bright, high-standing
exposing troughs and

Mosaic of
Mars Orbiter
Camera images



PSP_009656_2780



PSP_001334_2645

Finely layered deposits (NPLD)
exposed in a polar trough

PSP_001871_2670

MOC:

NASA/JPL/Malin
Space Science Systems

HiRISE images

NASA/JPL/University of Arizona

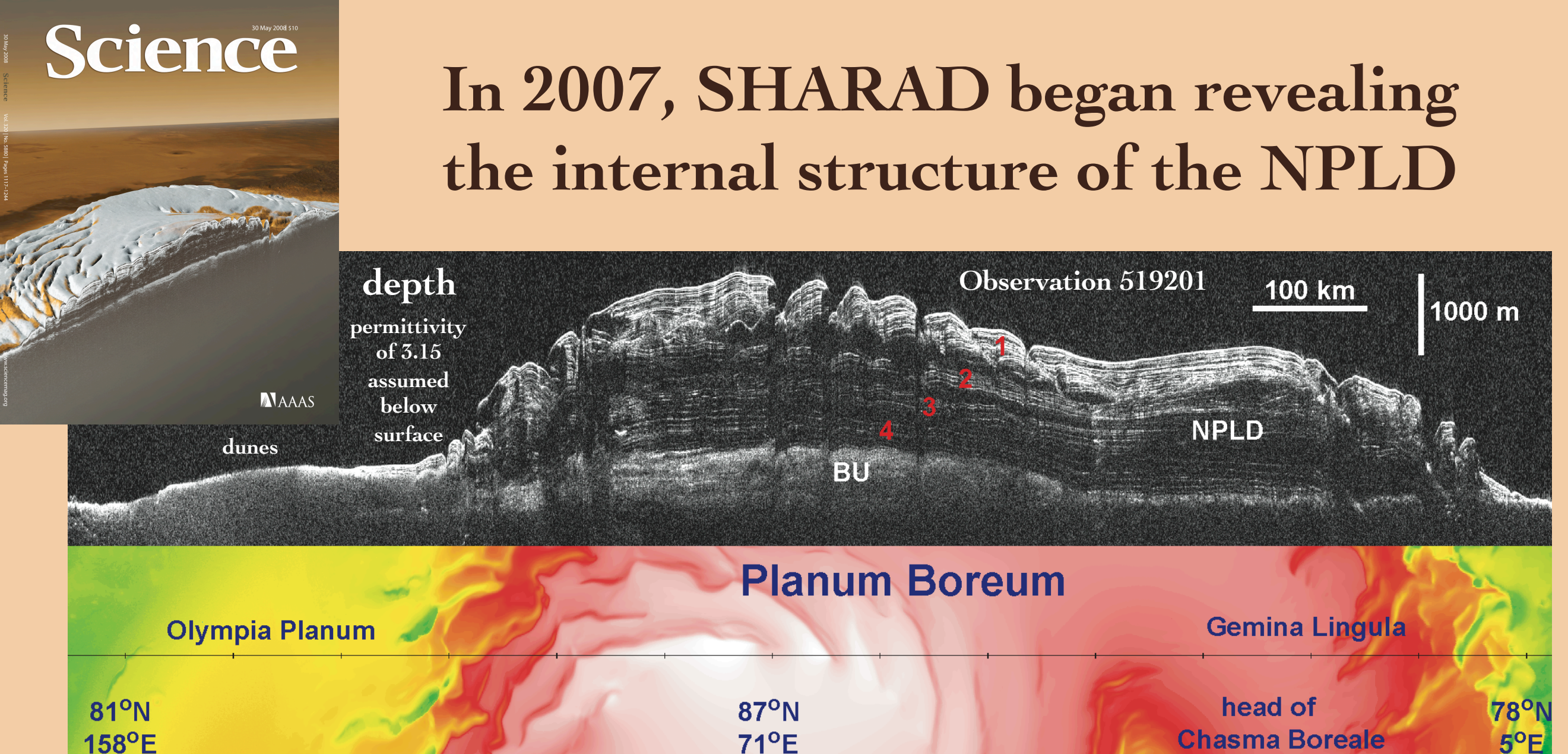
As MRO orbits Mars 12 times per day,
SHARAD reveals layering within the subsurface.



Animation credit: NASA/JPL-Caltech/University of Rome/SwRI

Orbit altitude: 255 to 320 km Wavelength: 15 m (~9-m vertical resolution in ice)
Transmitted sweep: 15 to 25 MHz Lateral resolution: 3 to 6 km (0.3 to 1 km inline with SAR)

In 2007, SHARAD began revealing the internal structure of the NPLD

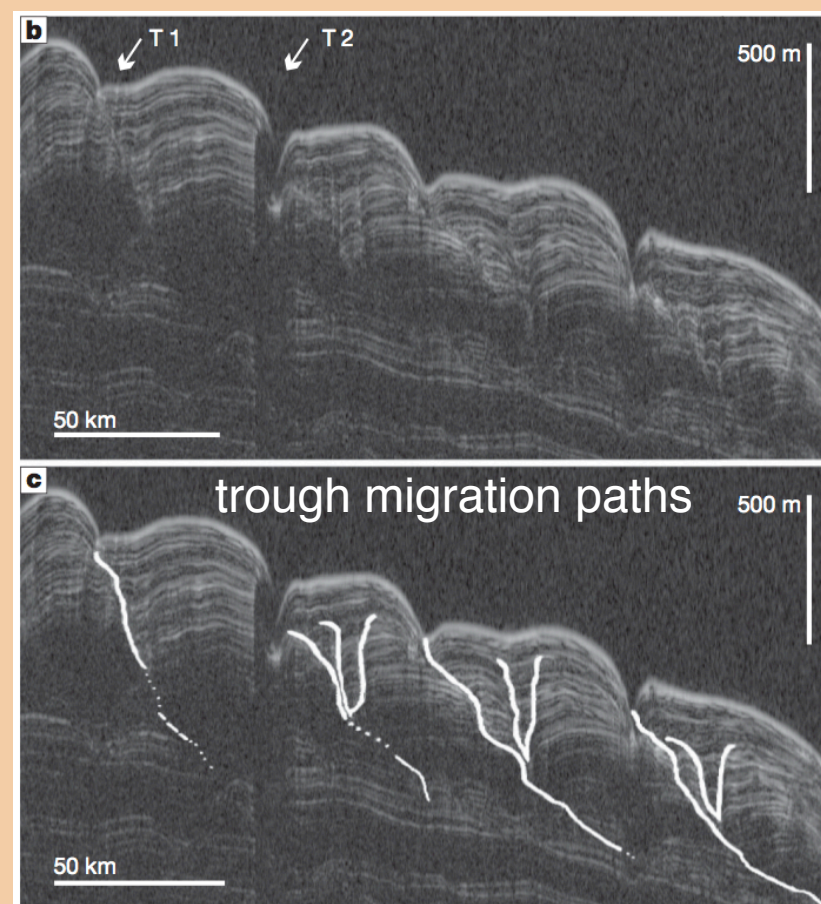
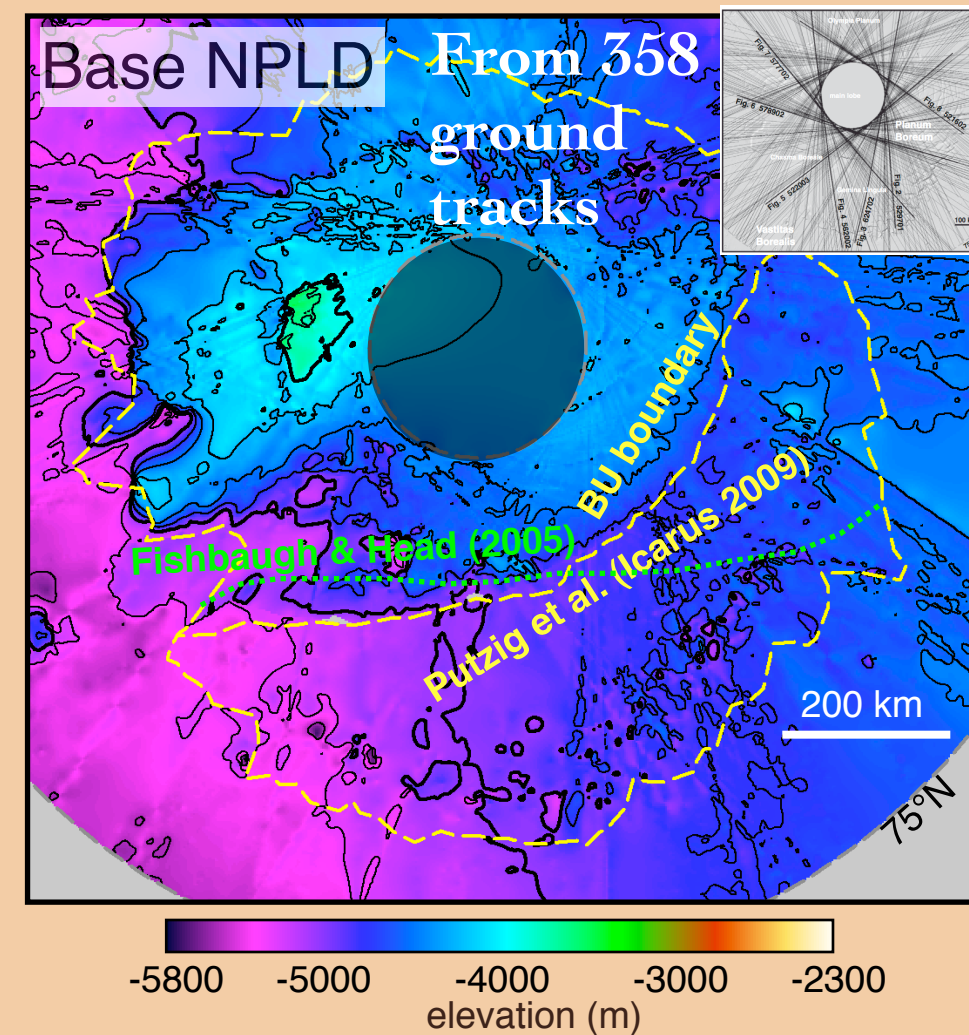


- Strong basal returns imply relatively pure ice (<~5% lithics)
- NPLD packet structure likely related to climate; age ~ 4.2 Ma
- Older basal unit rarely layered, missing below >1/3 of NPLD
- Flat basal boundary has ~ 0 flexure \Rightarrow very low heat flow

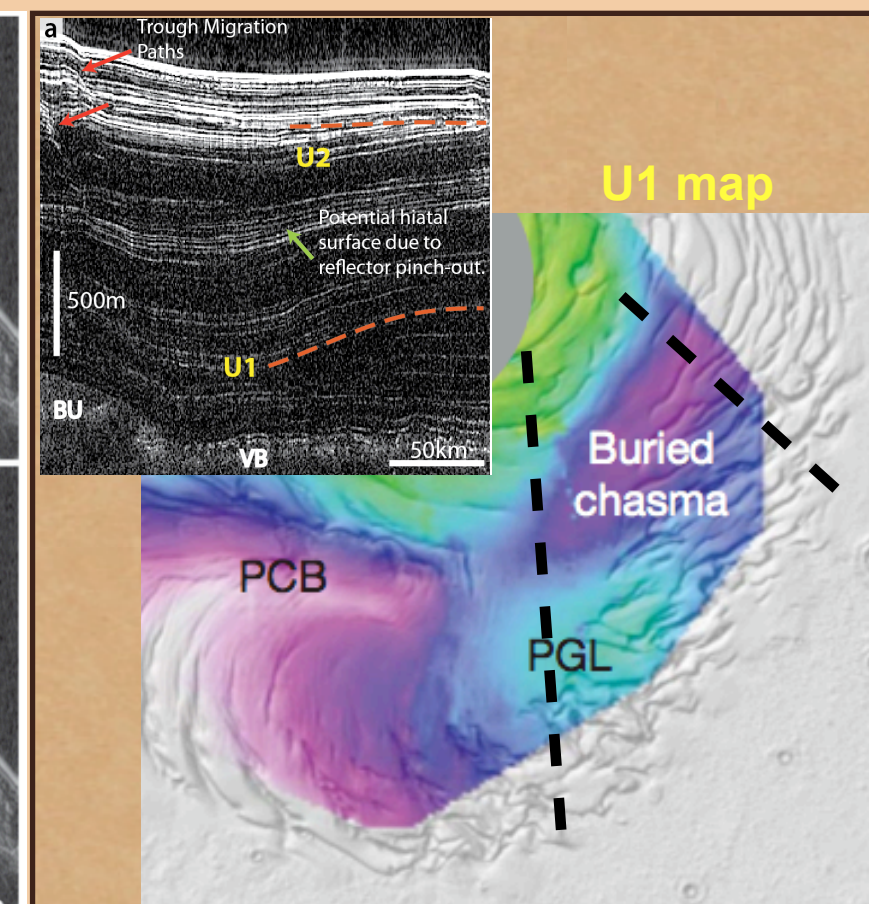
Phillips et al. (Science 2008); Putzig et al. (Icarus 2009)

A growing grid of 2-D radar data is yielding important new science:

- Revised BU boundary led to a change of view on Chasma Boreale from an erosional to a constructional feature.
- Spiral troughs are shown to be wind-driven icy bedforms that are actively migrating poleward.
- Mapping internal unconformities revealed a buried chasma with no surface expression.



Smith and Holt (Nature 2010)



Holt et al. (Nature 2010)

Limitations of conventional 2-D radar analysis

- **Clutter:** off-nadir returns that interfere with or can be mistaken for nadir-surface or subsurface returns.
- No returns are seen from features at nadir that slope away from the radar.
- Topographic models can be used to simulate surface returns and clutter. Corresponding features in data may then be dismissed as “noise”.
- However, subsurface clutter and off-nadir surface features not sampled in topographic data will not appear in synthetics.

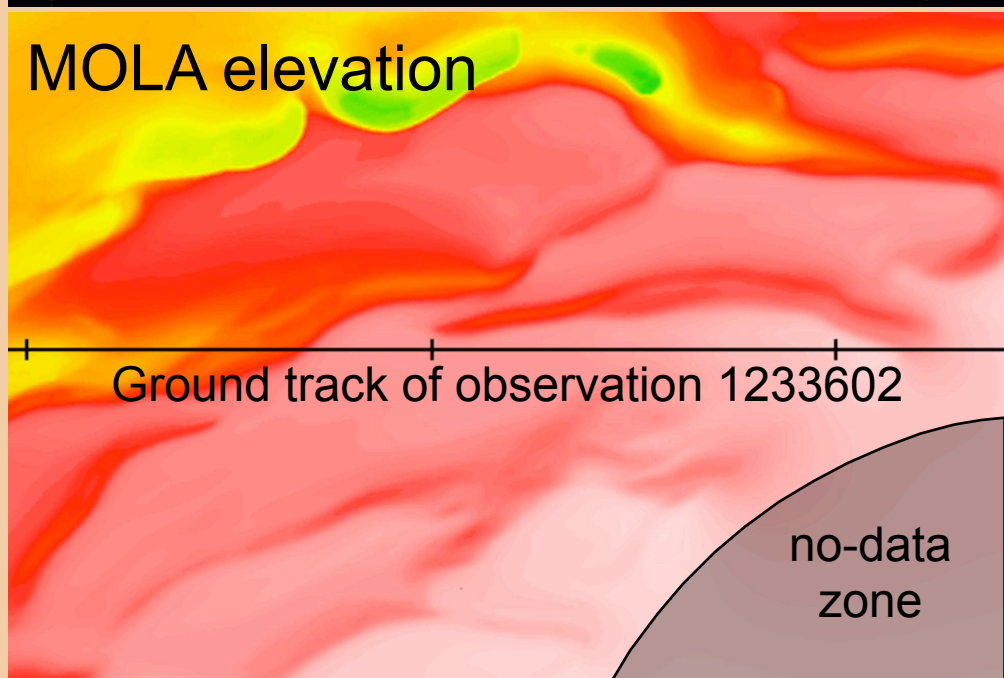
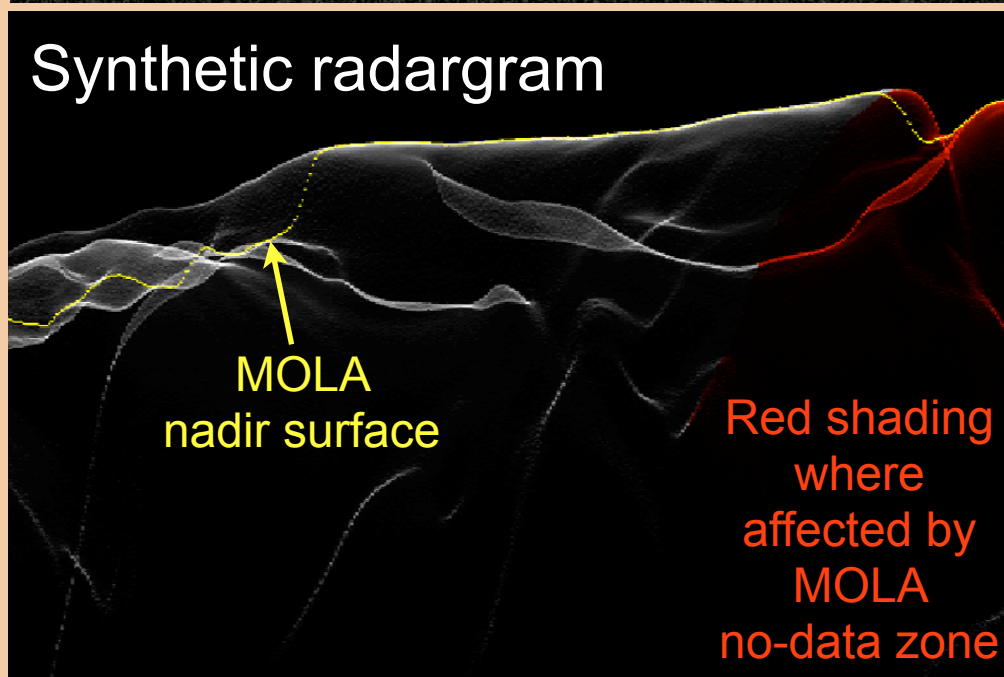
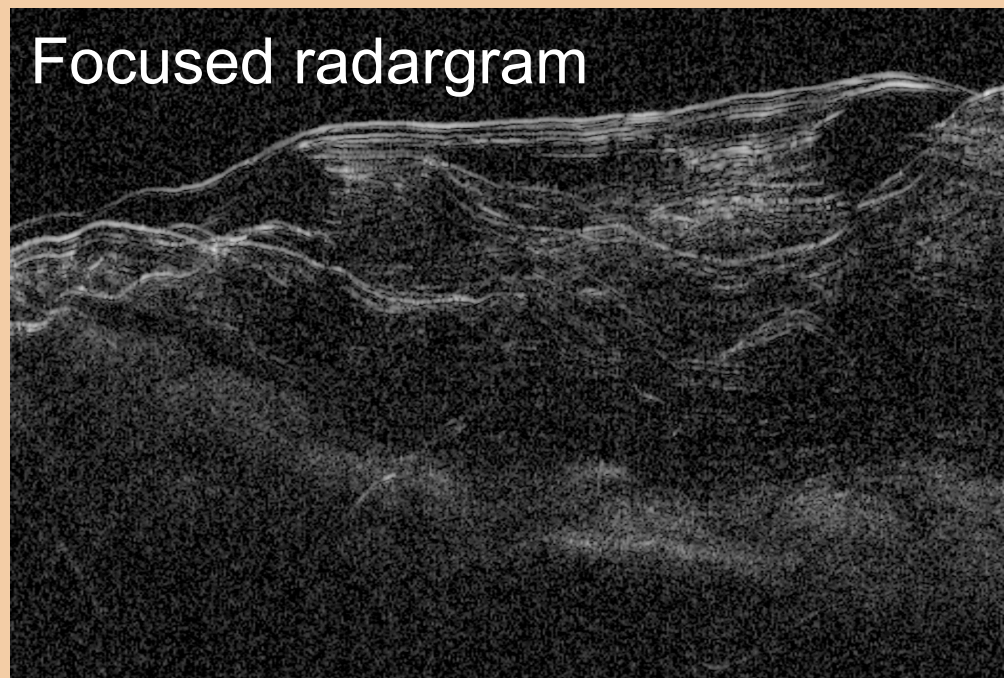
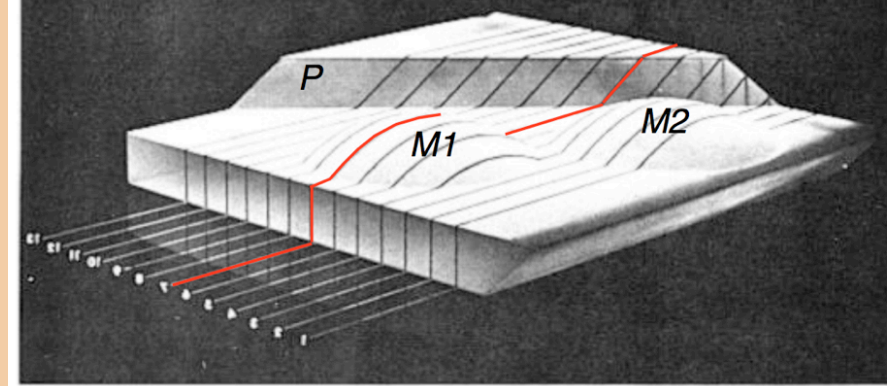
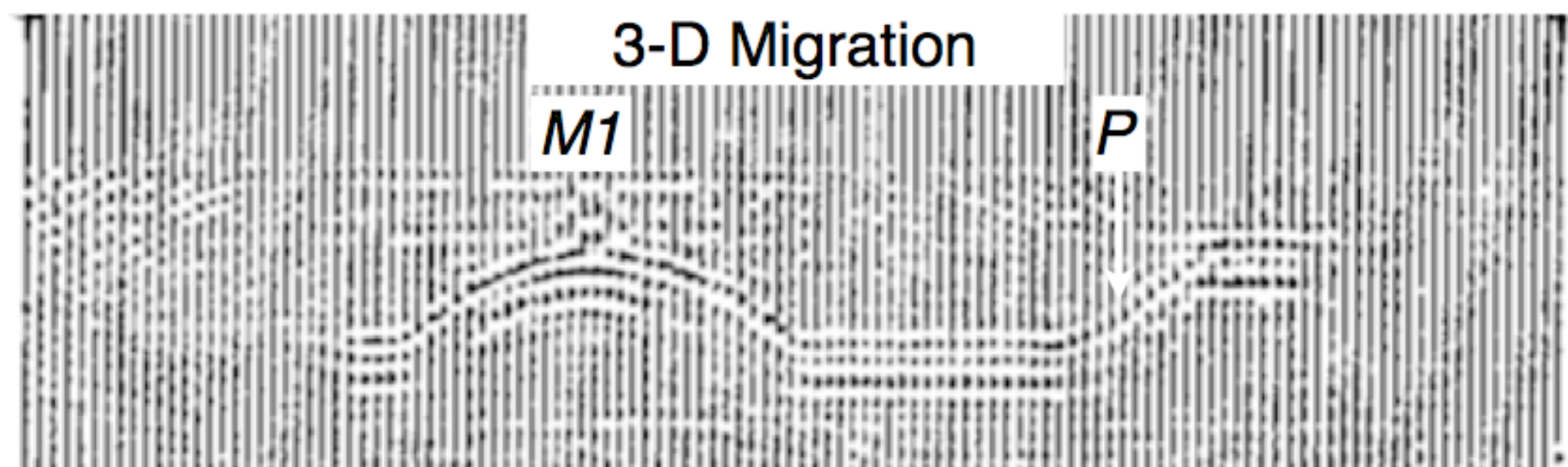
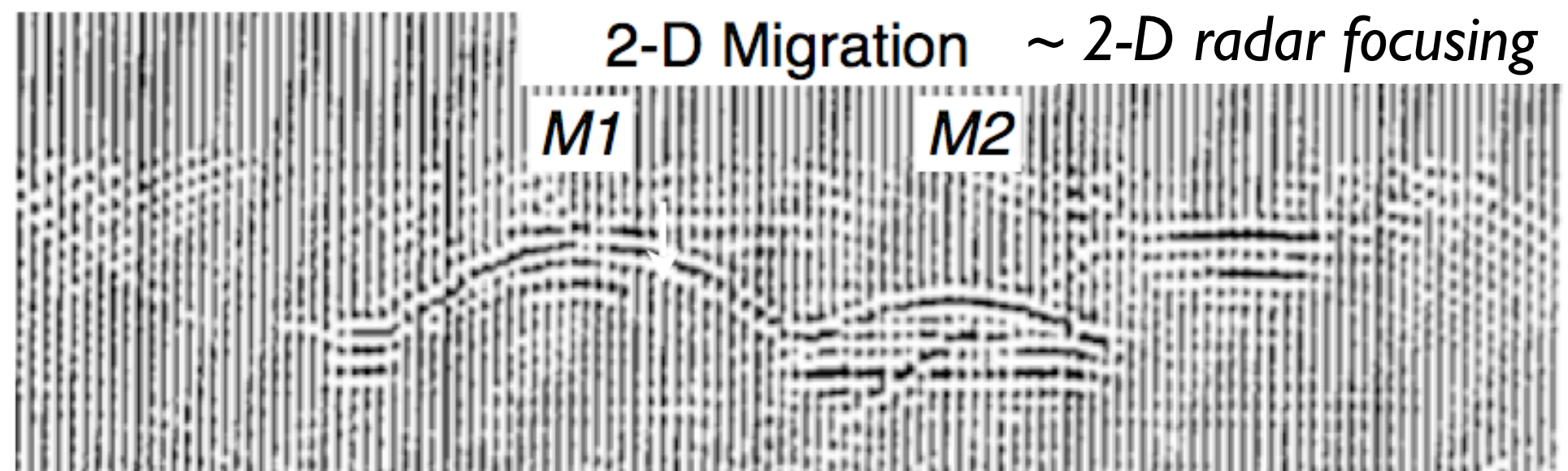
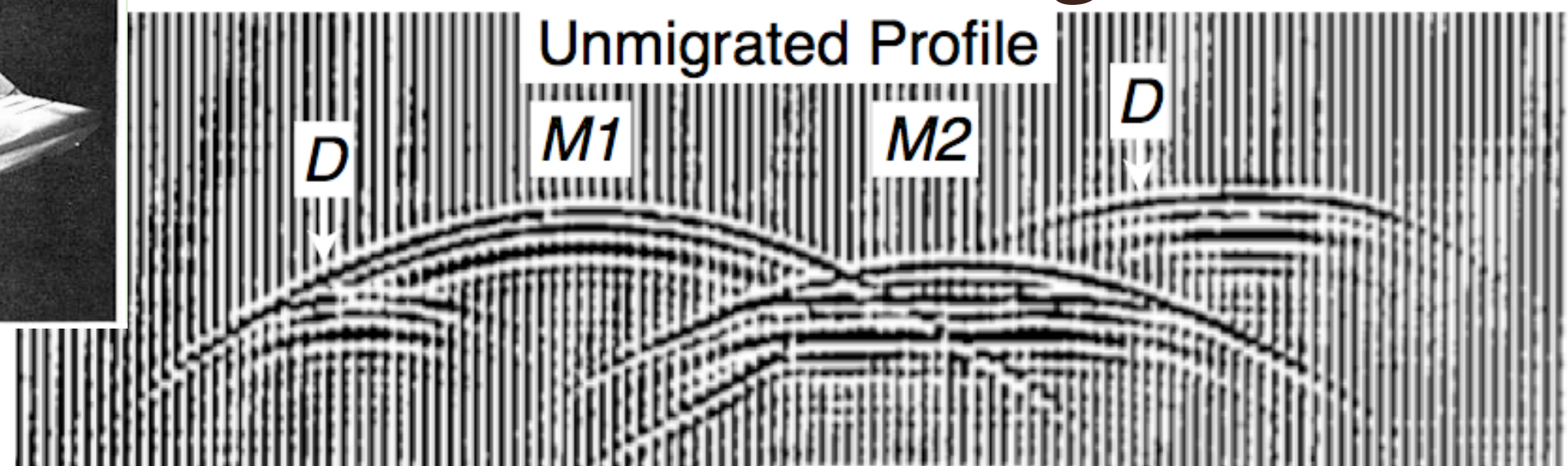


Photo of the Physical 3-D Model



Solution: 3-D migration



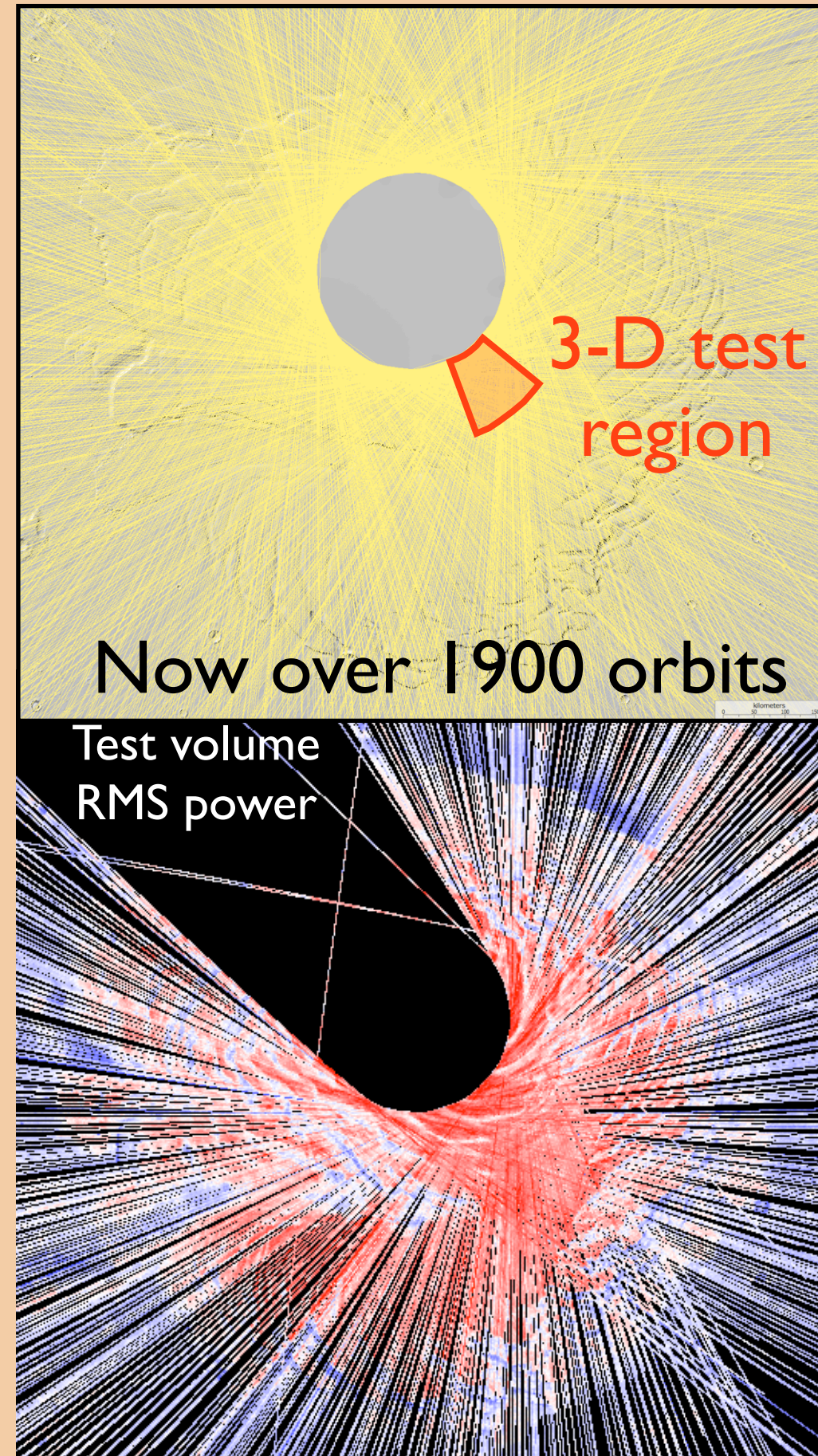
3-D migration will collapse diffractions (D) and reposition out-of-plane returns (M2) to their source locations.

Energy from adjacent profiles will be restored, thereby imaging features oriented obliquely to the profile (P).

The SHARAD 3-D volume

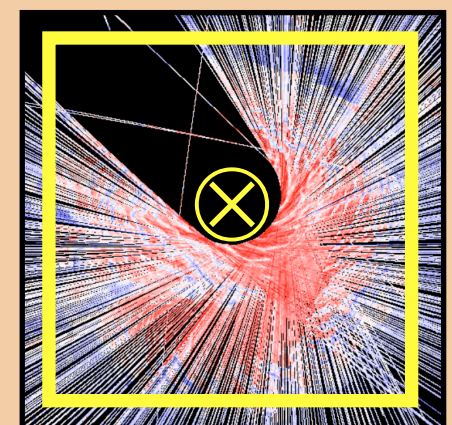
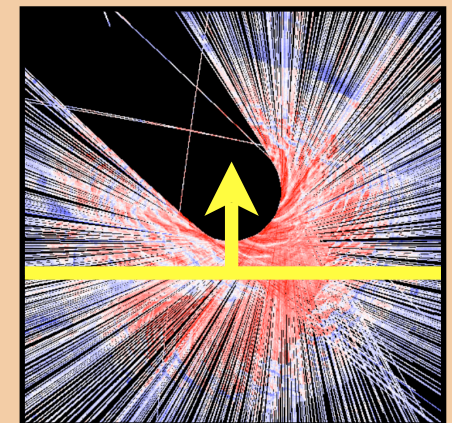
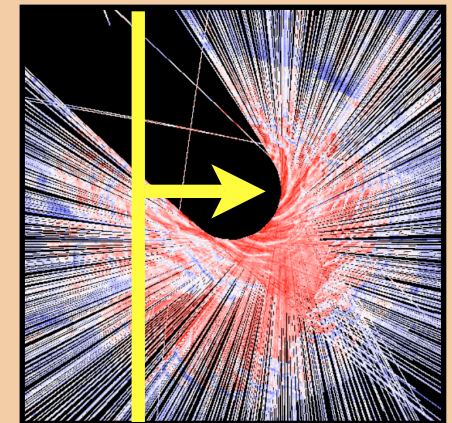
- We reference the 2-D radargrams to a common datum* and bin them into a rectilinear grid.
- Our 3-D grid covers 1500×1500 km with **9 million bins** of 500×500 m.
- For **binning and migration tests**, we selected 540 radargrams crossing the Planum Boreum saddle region.

* *Non-trivial step due to variable ionospheric delays*



Movie of the test volume: (prior to 3-D migration)

- First sequence shows **inline** profiles scanning from left to right across the 3-D grid.
- Second sequence shows **crossline** profiles scanning from bottom to top over the 3-D grid.
- Third sequence shows **timeslice** views scanning in delay time (into the subsurface).

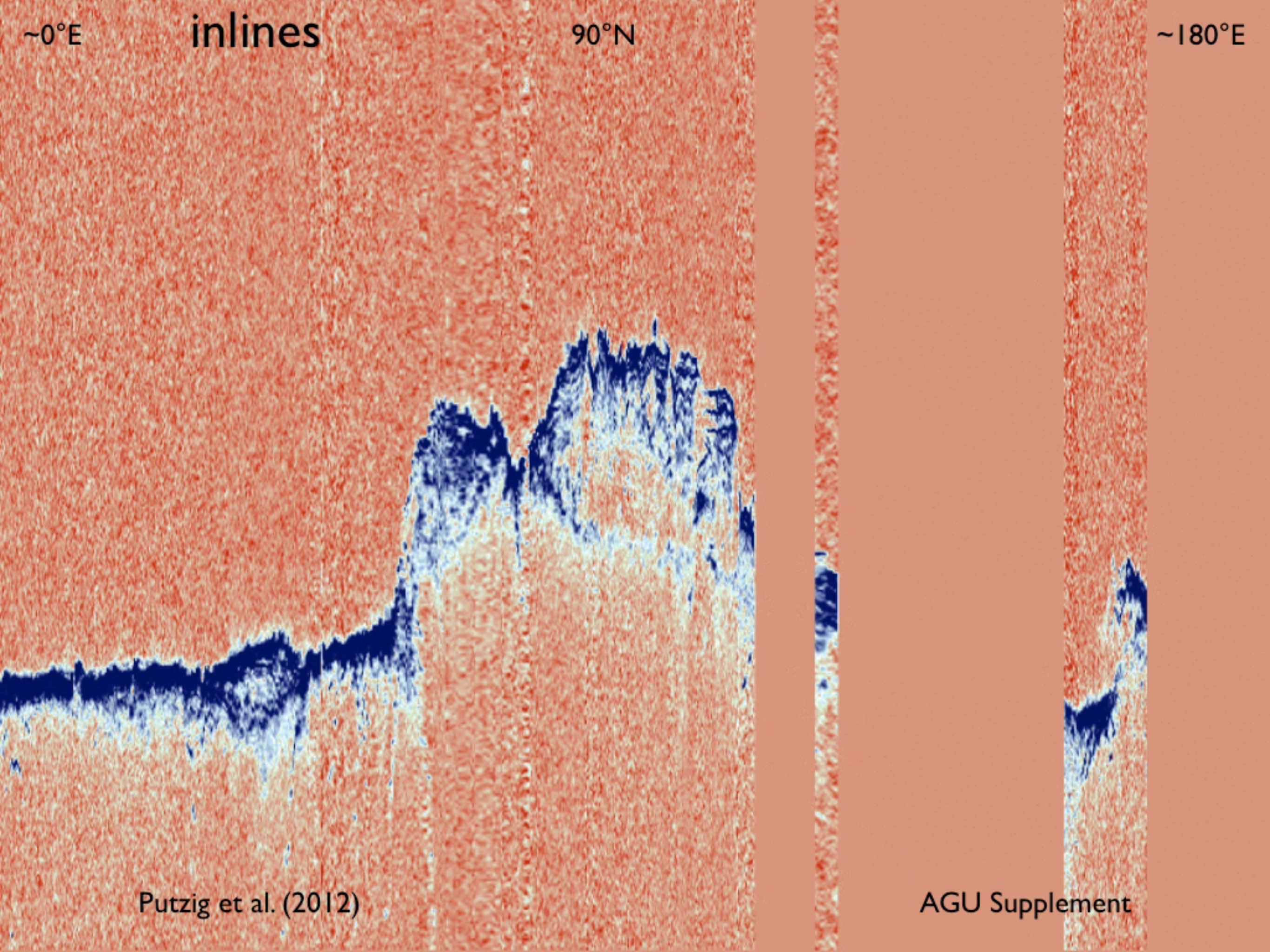


~0°E

inlines

90°N

~180°E

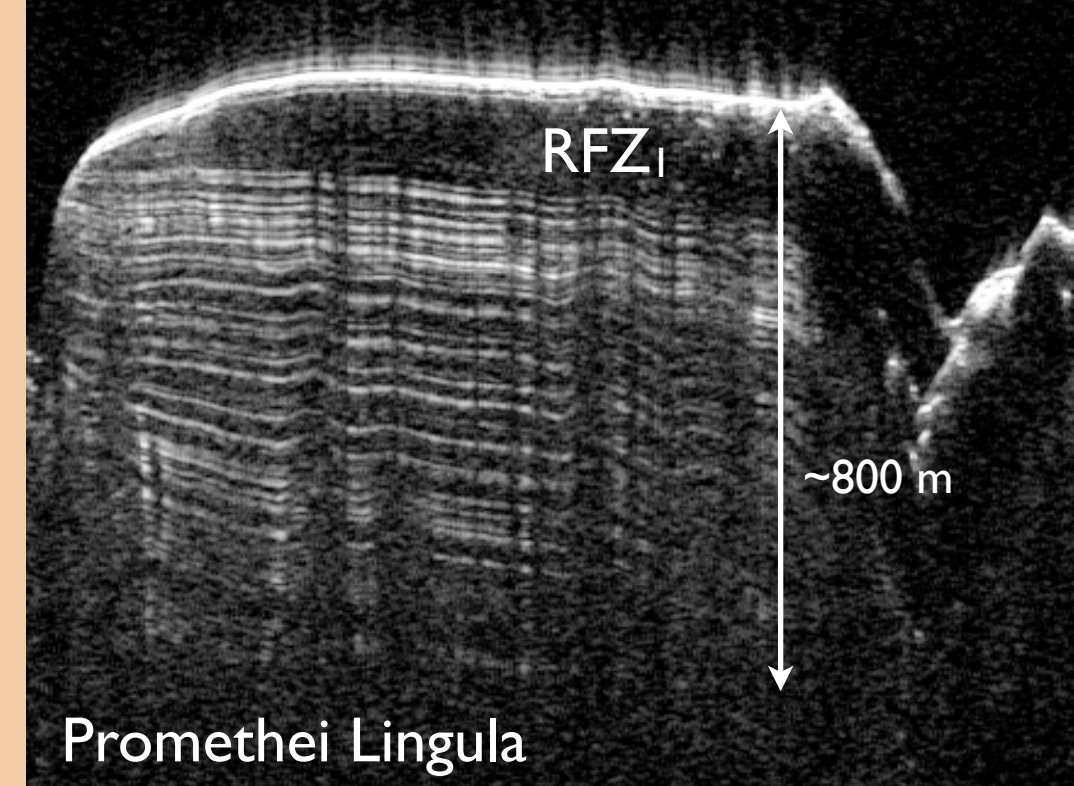


Putzig et al. (2012)

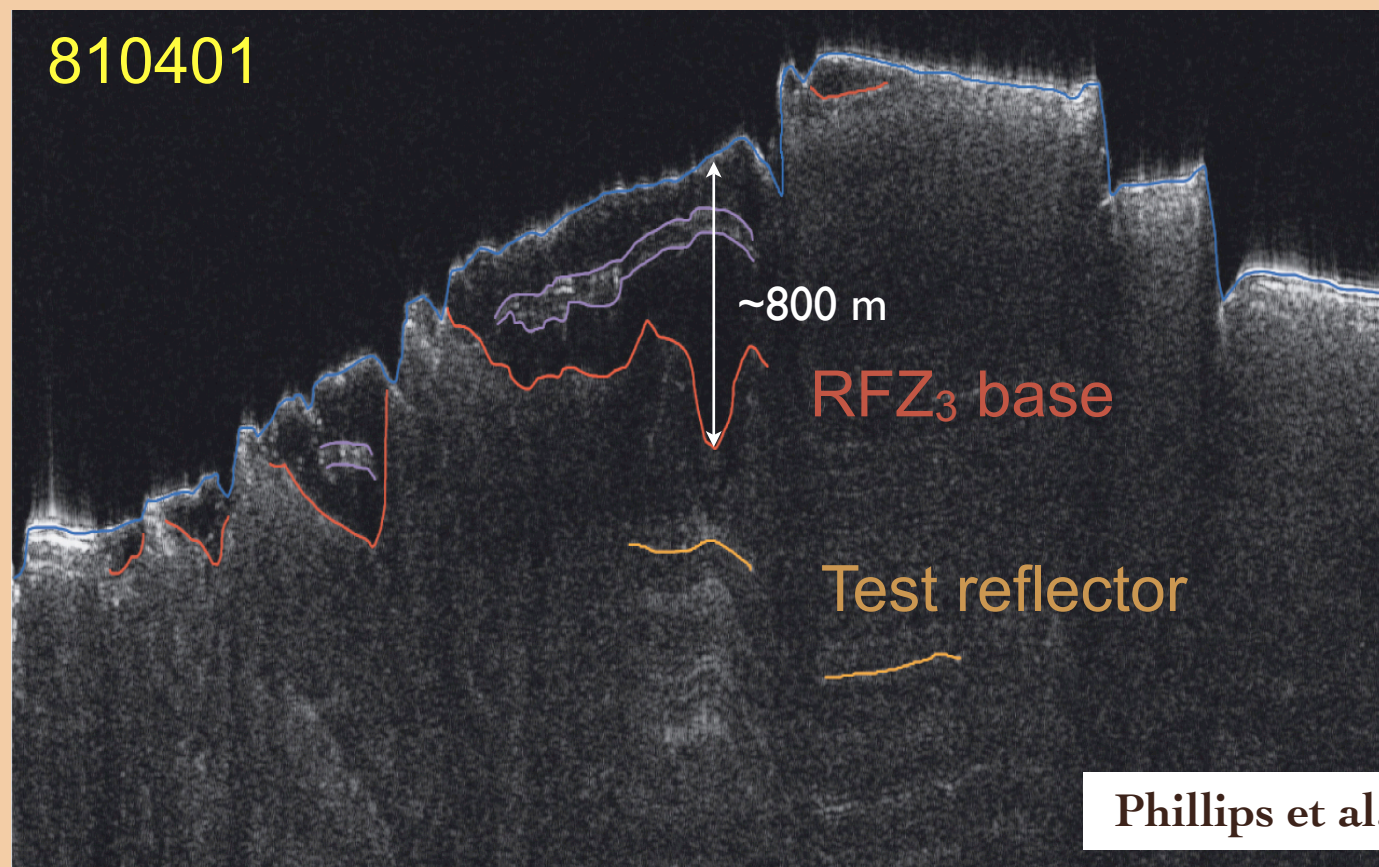
AGU Supplement

What about the south?

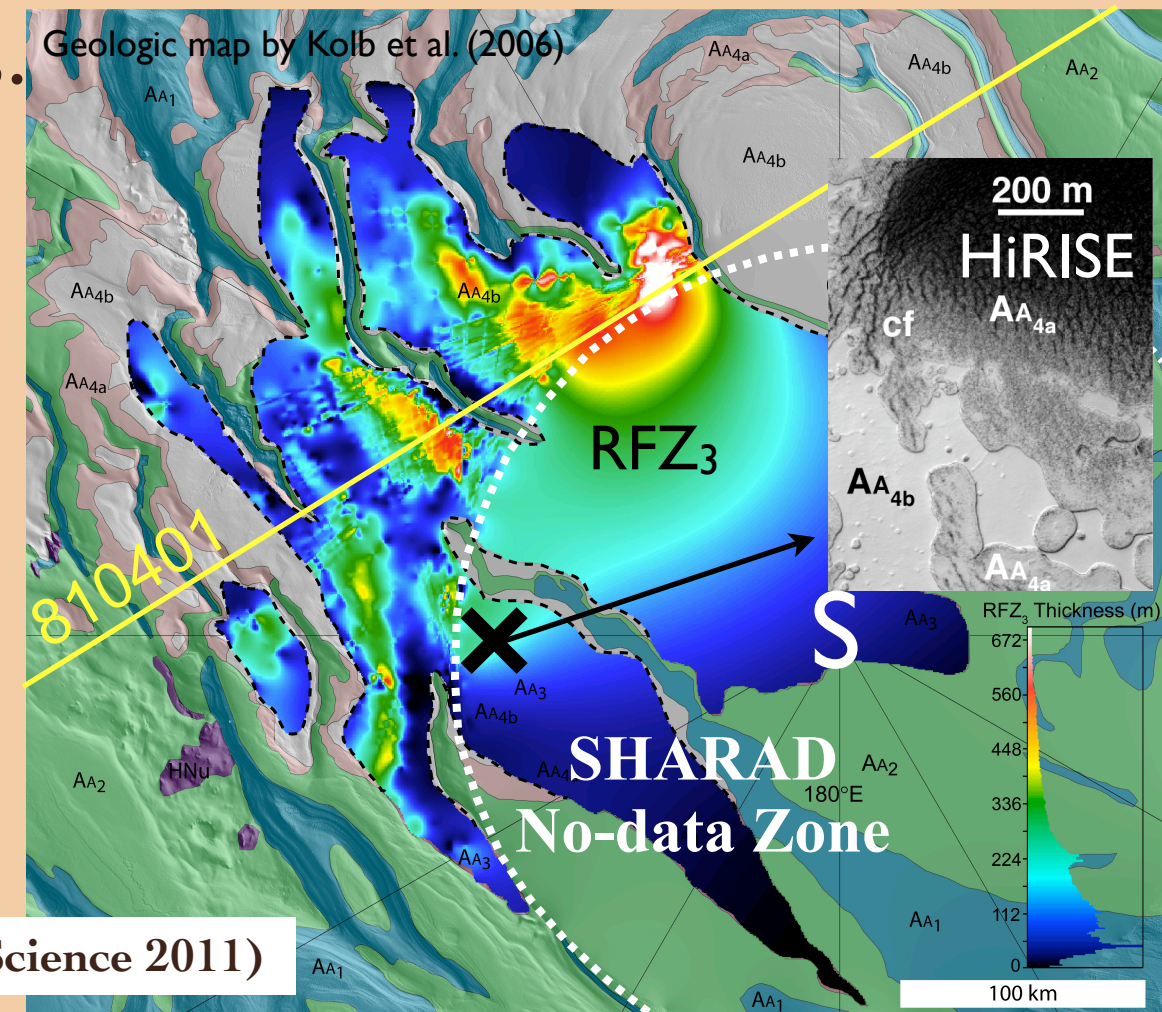
- Layering is discontinuous and often truncated just below the surface
⇒ likely much older than NPLD.
- Odd reflection-free zones (RFZs) occur near surface. Near the pole, RFZ₃ has dielectric properties consistent with CO₂ ice and maps to a unit with CO₂ sublimation features.



4–5 mbar of CO₂!

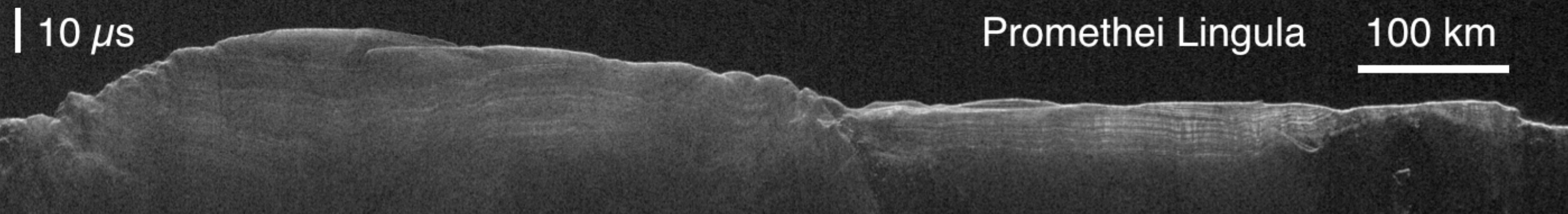


Phillips et al. (Science 2011)

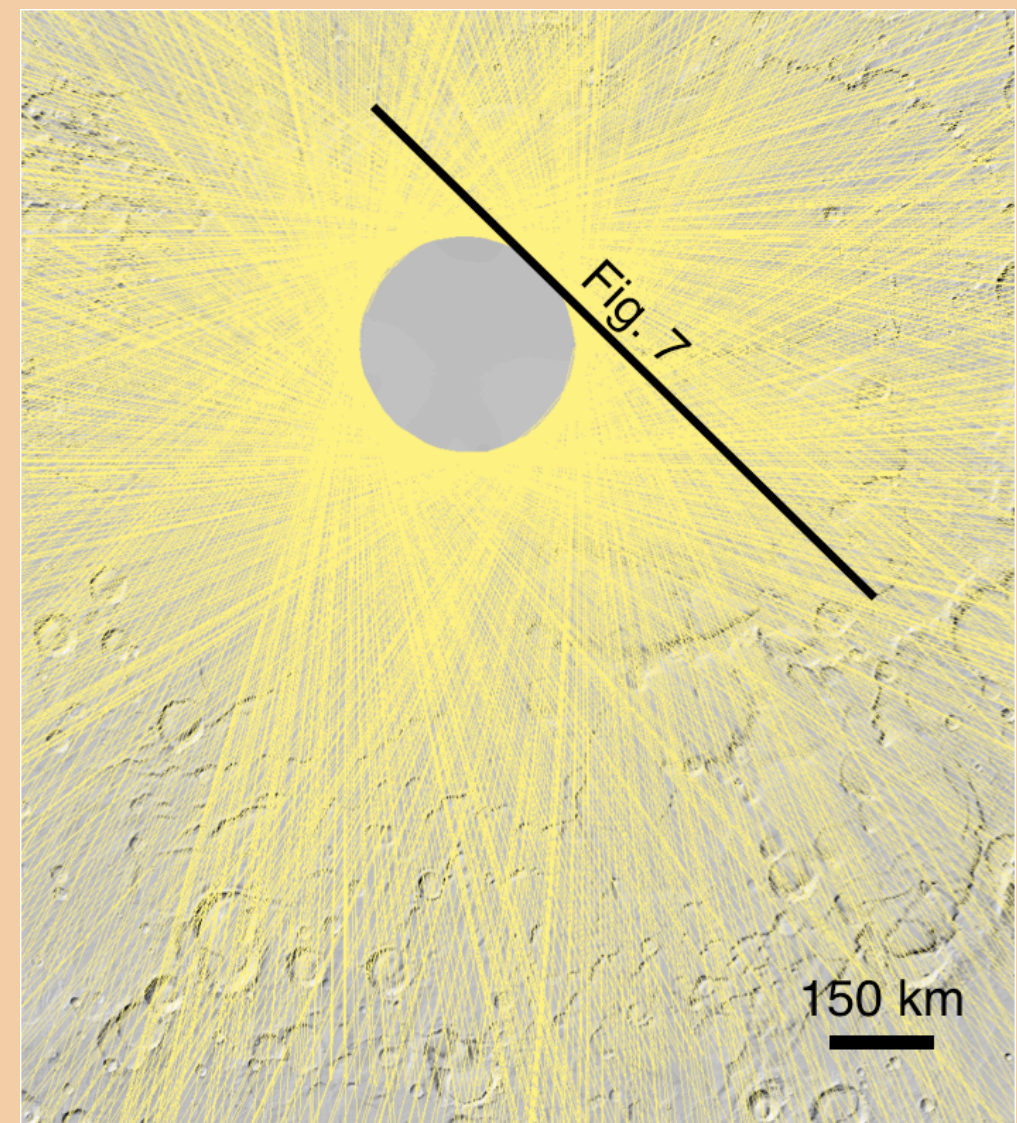


3-D for the Planum Australe?

Please fund our MDAP



- SNR improvements and geometric corrections of 3-D migration may be critical for:
 - ▶ mapping water-ice layers
 - ▶ assessing RFZs 1, 2, & 4



Conclusions

- Analysis of 2-D SHARAD radar data has yielded a wealth of discoveries in the Martian polar ices.
- 3-D binning and migration processing promise to add a new level of clarity to this rich volume of data.
- Further improvements will shed new light on the nature and timing of the polar deposits at both poles.

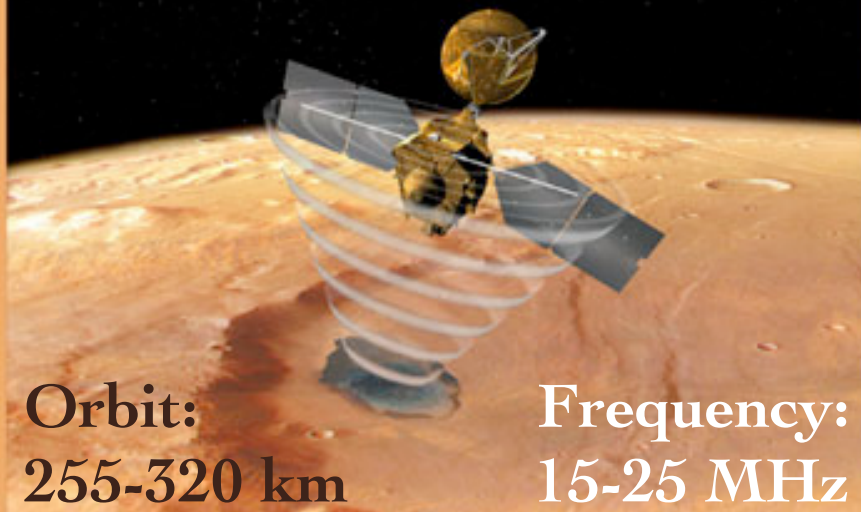
Thanks to...

- SeisWare for software access.
- NASA, ASI (Italian Space Agency), the MRO Project, and the SHARAD Instrument Team.
- NASA's Mars Data Analysis Program for funding this work.

extra slides

MRO's Shallow Radar sounder

SHARAD



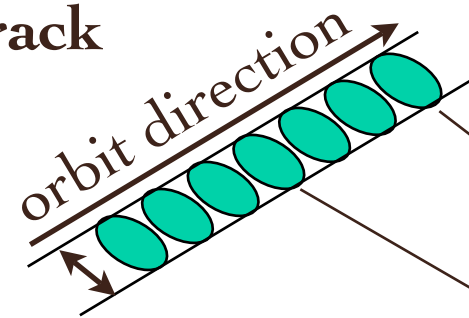
Orbit:
255-320 km

Frequency:
15-25 MHz

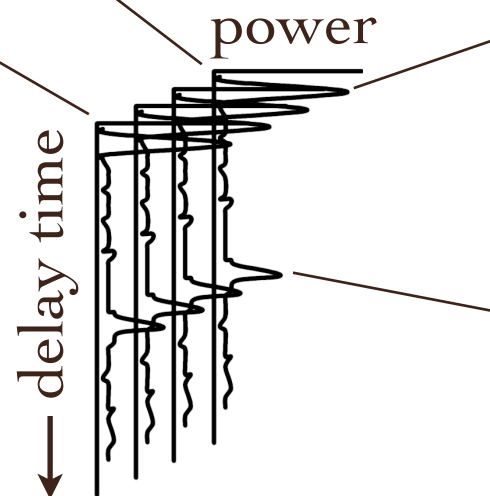
Primary Objective

Map subsurface dielectric interfaces and interpret them in terms of the occurrence and distribution of expected materials, including rock, regolith, water, and ice.

Ground Track



Individual Echoes



Resolution

Cross-track: 3-6 km

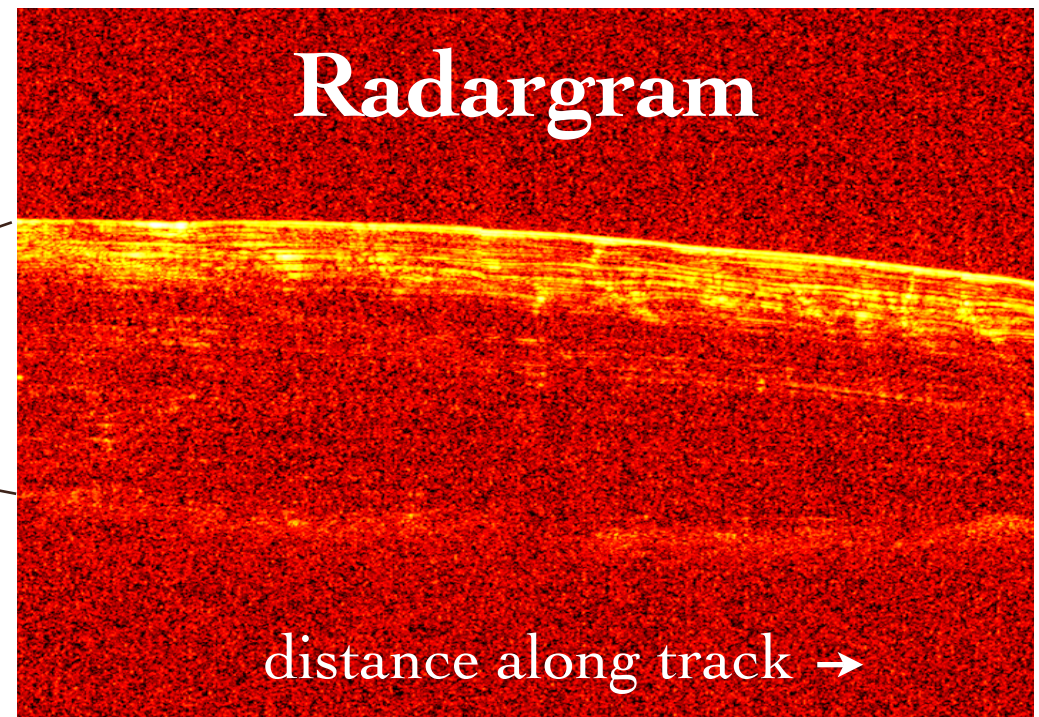
Along-track: 0.3-1 km*

Range: $15 \epsilon_r^{-1/2}$ m

≈ 8 m in H₂O ice

≈ 10 m in CO₂ ice

Radargram



* Along-track resolution is improved using synthetic aperture radar (SAR) processing techniques

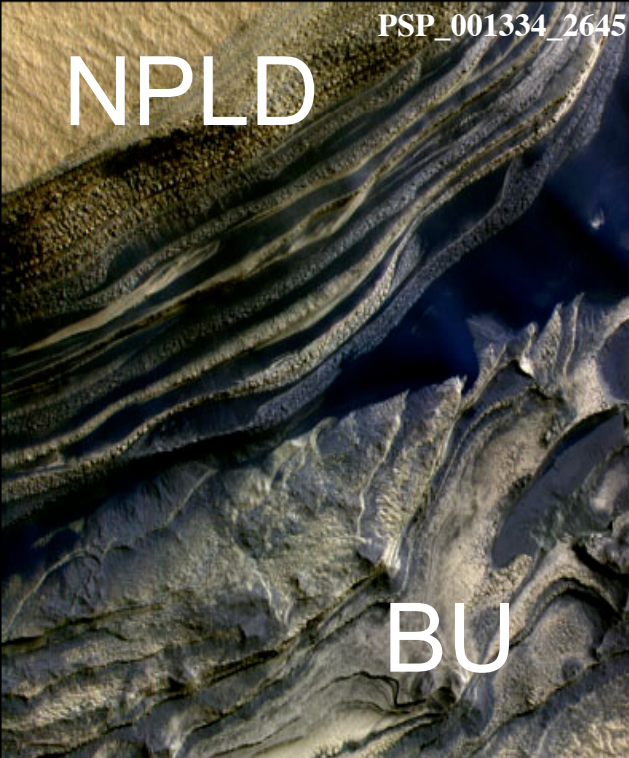
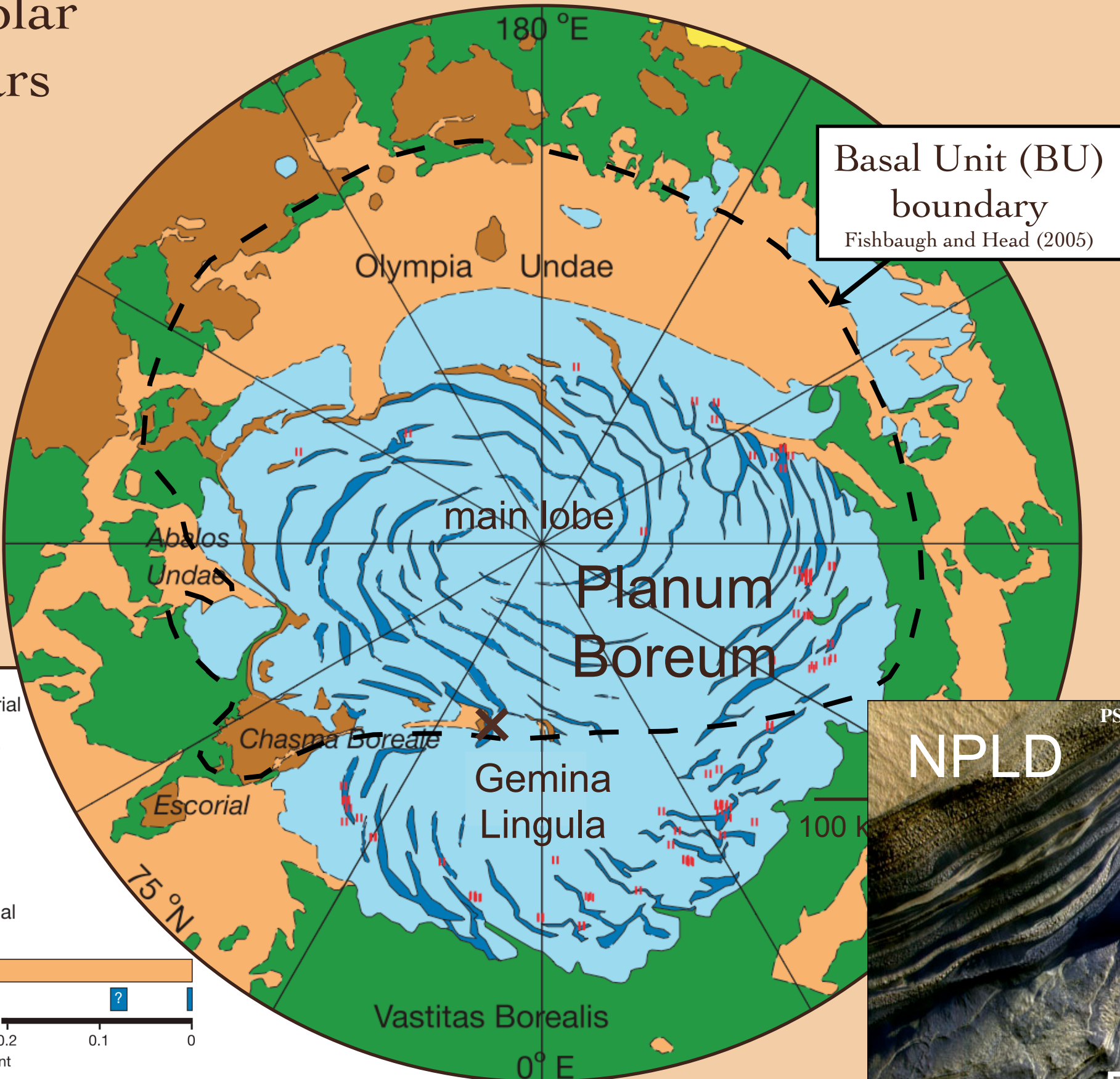
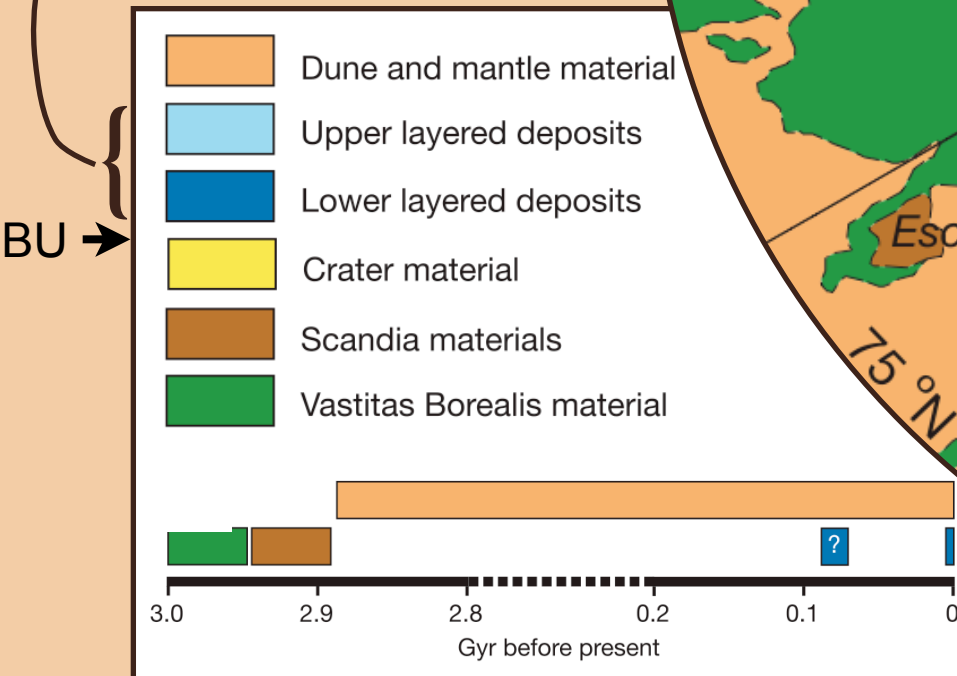
The North Polar Region of Mars

Geologic map

Tanaka (2005)
Nature 437, 991

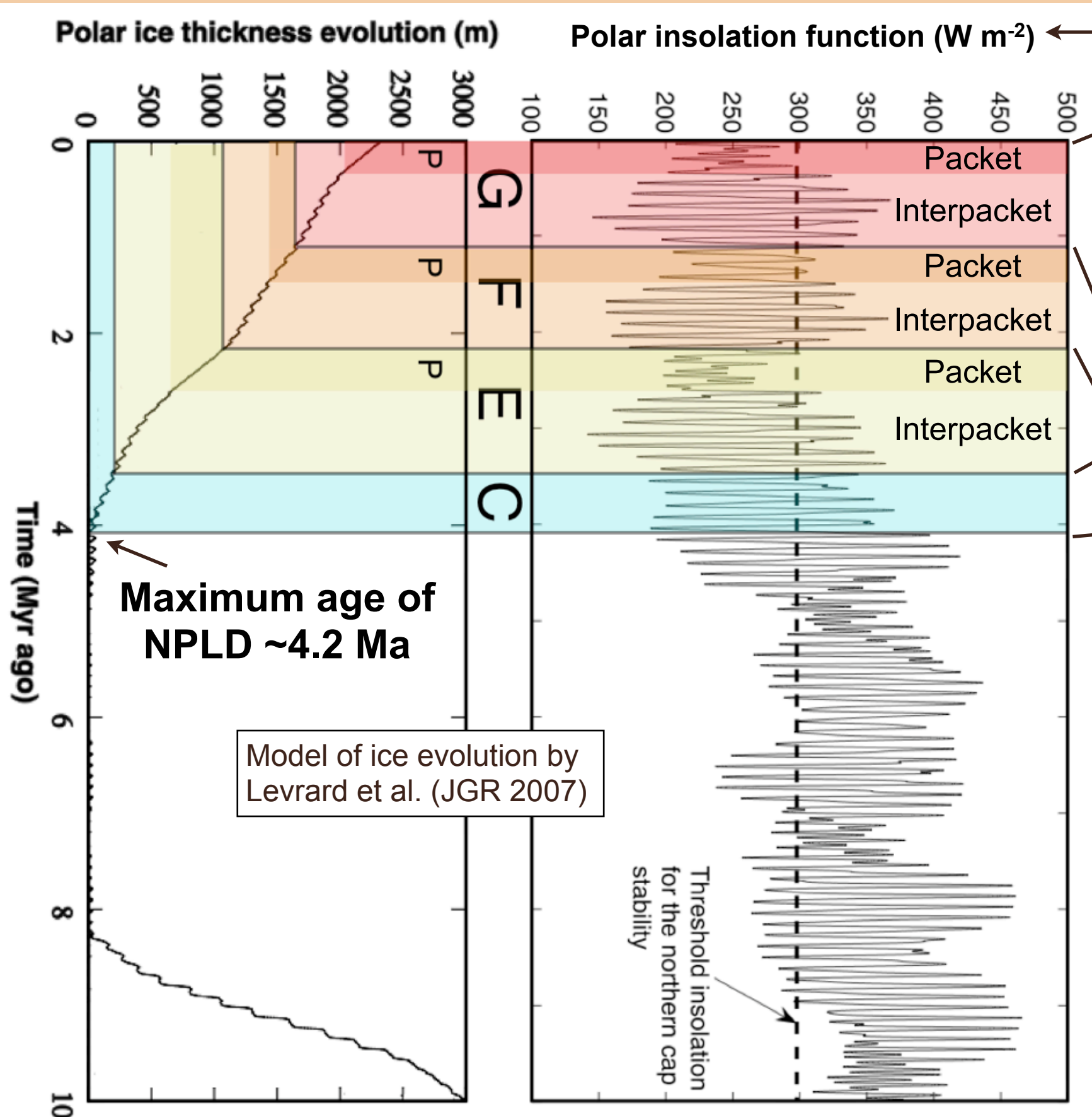
USGS
nomenclature

North Polar Layered Deposits (NPLD)

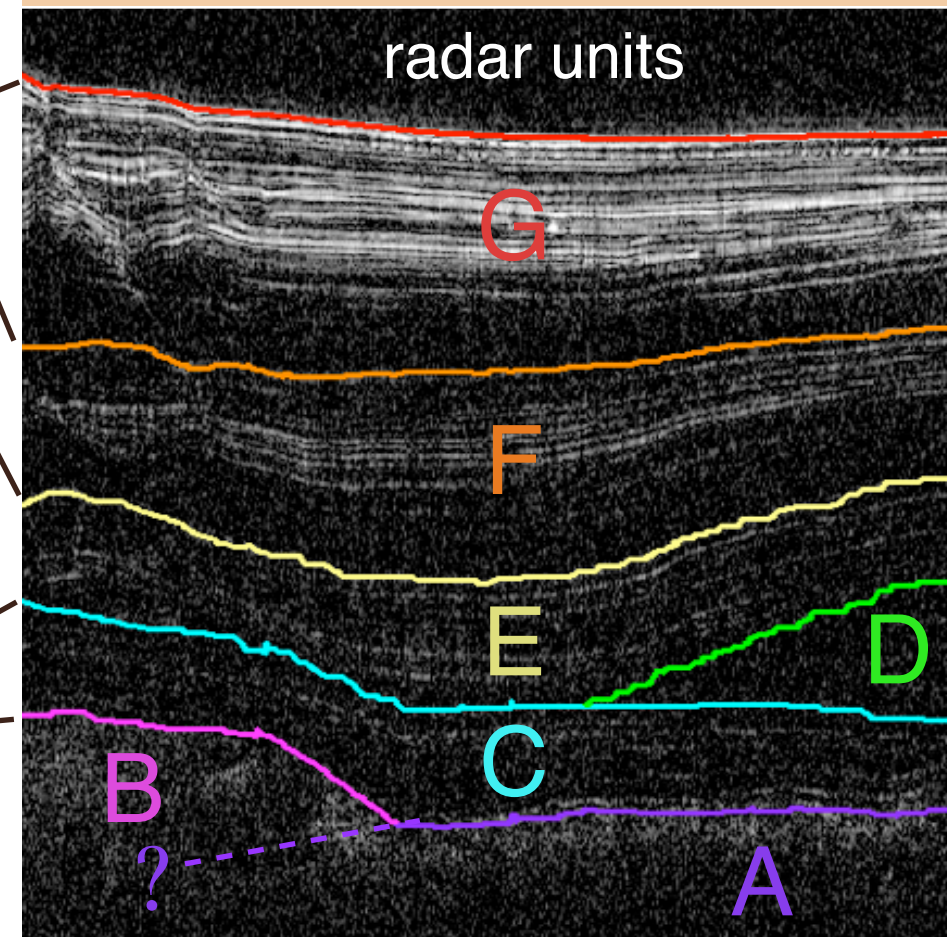


HiRISE image credit:
NASA/JPL/University of Arizona

Correlation of radar units to climate cycles



Martian orbit is well-constrained from -20 Ma to +10 Ma (Laskar et al., 2004)



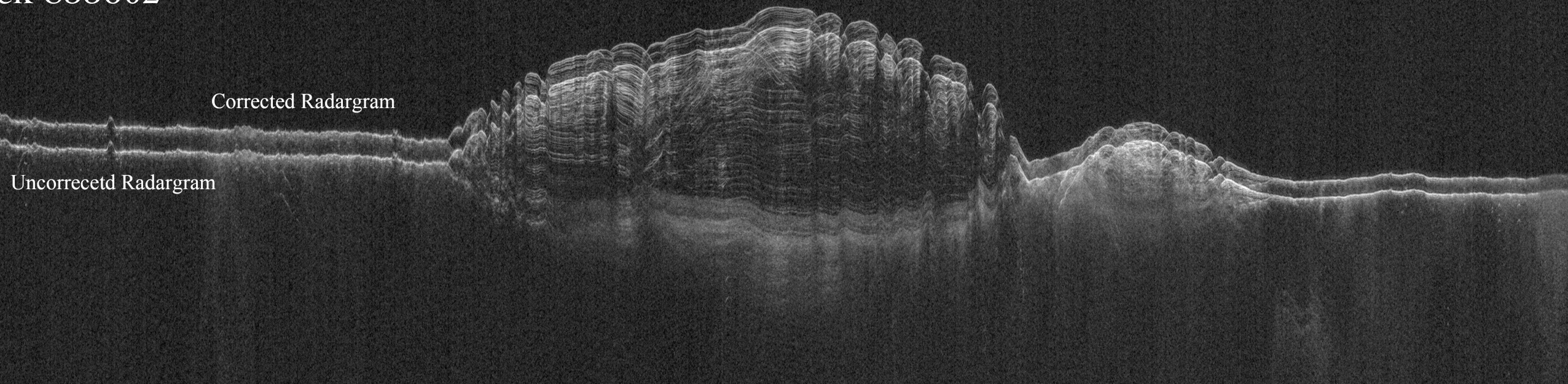
Strongly layered packets may correspond to periods with low amplitude in the variation of insolation — the present era, ~1.3 Ma, and ~2.4 Ma.

Putzig et al. (Icarus 2009)

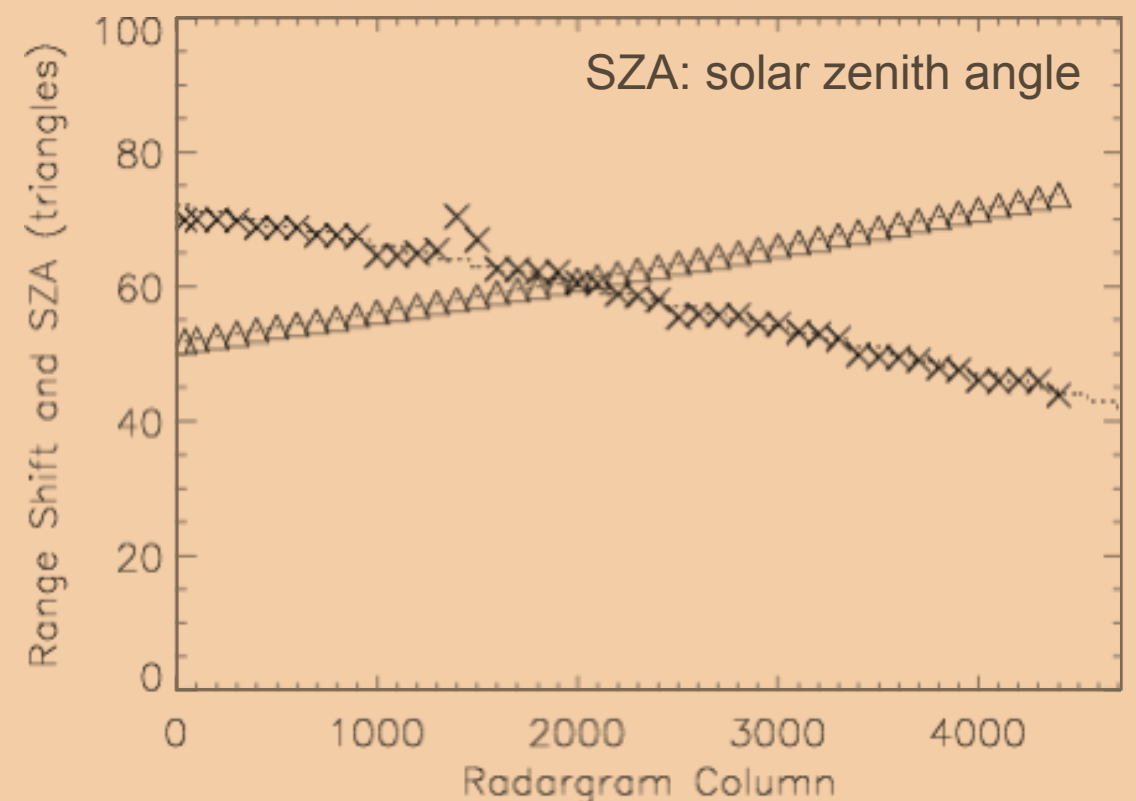
Challenge to 3-D Processing

- Comparison of crossing 2-D radargrams at their intersections revealed variable shifts in range (delay time), most pronounced in data acquired on the daylit side of the planet:

Track 855602



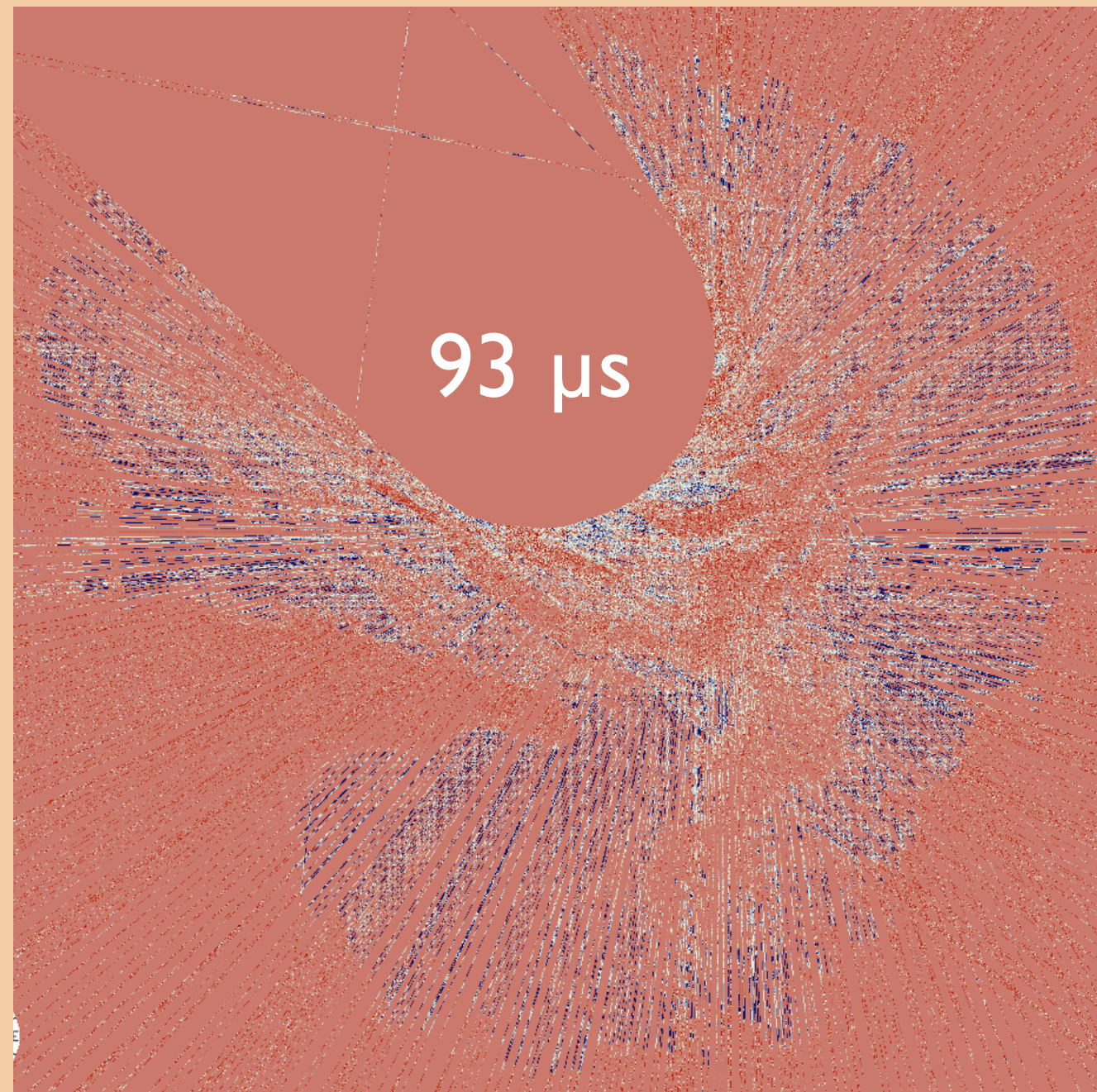
- The delay is attributable to the Martian ionosphere, and can be approximated by a linear fit to the phase distortion derived in the pre-processing (× symbols at right).
- Correction greatly reduces mismatches in vertical positioning and is critical for proper 3-D migration.



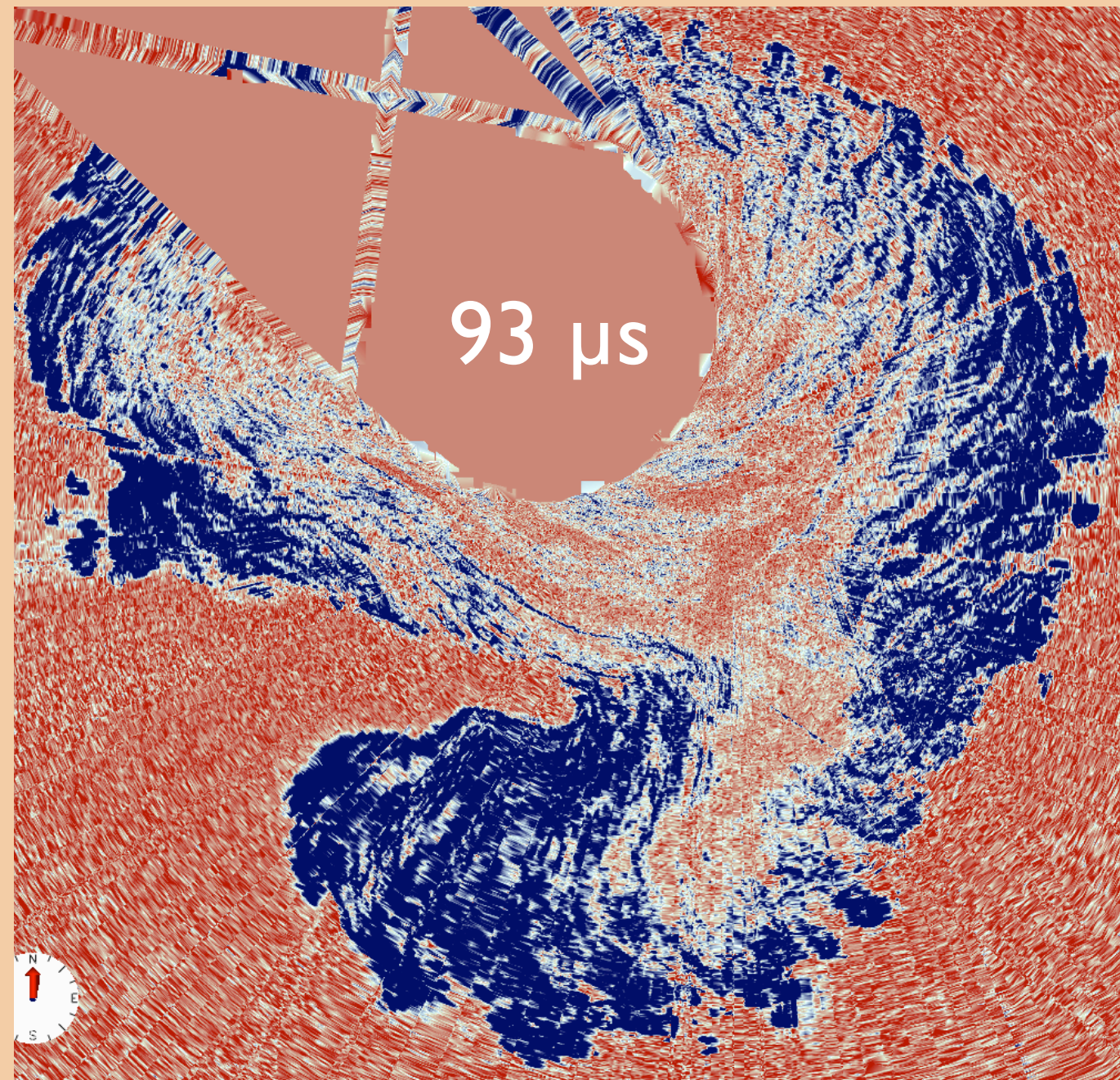
3-D Binning Process

Spatial Delauney triangulation, bilinear interpolation, and amplitude regularization in 9×10^6 bins of $500 \text{ m} \times 500 \text{ m}$.

Timeslice prior to interpolation and regularization



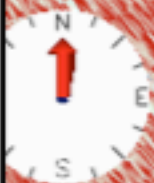
Timeslice after interpolation and regularization



Uninterpolated

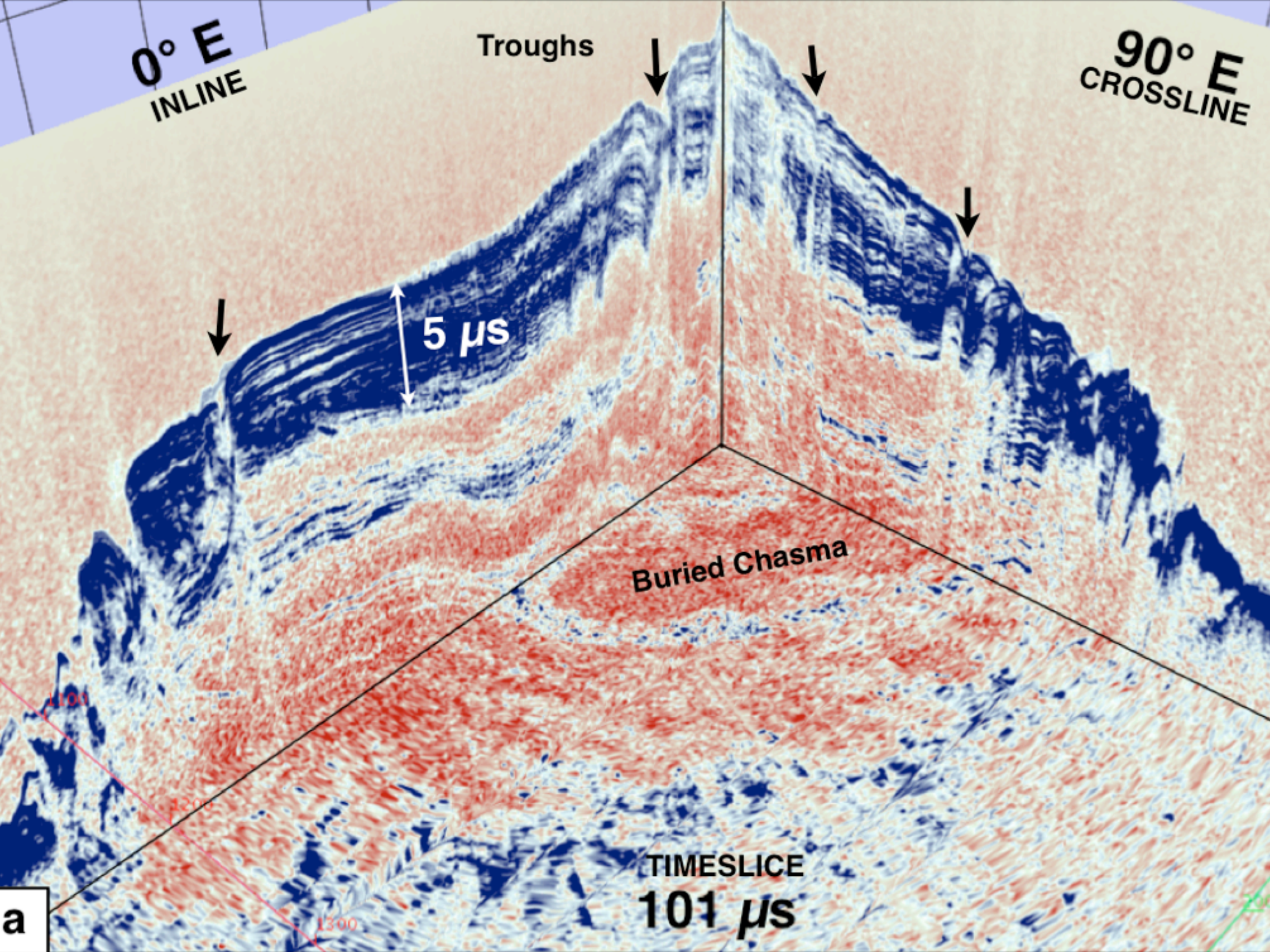
**Trough
migration
paths**

CB



TIMESLICE

93 μ s



a

